

Layered Convection and Convective Layers in Jupiter and other gas giant planets

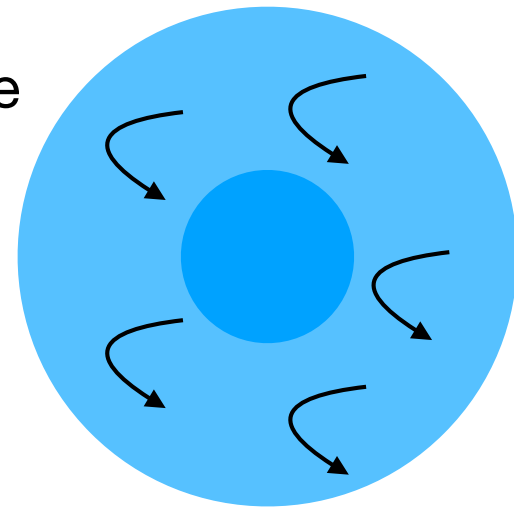
Andrew Cumming
McGill University



Motivation

- Composition gradients are a natural outcome of planet formation because heavy elements condense and are delivered separately (pebbles, planetesimals)

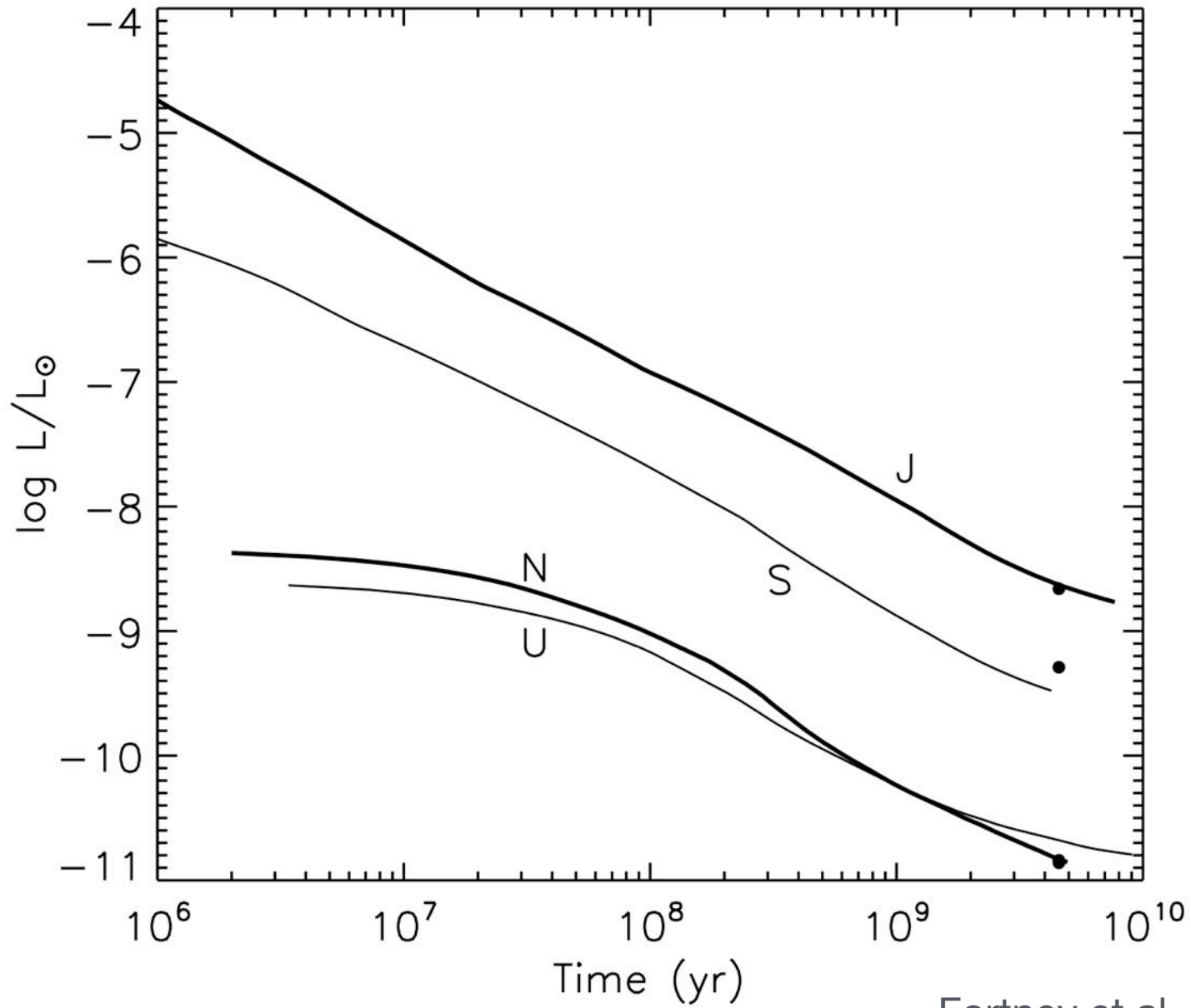
- Core accretion => core-envelope structure



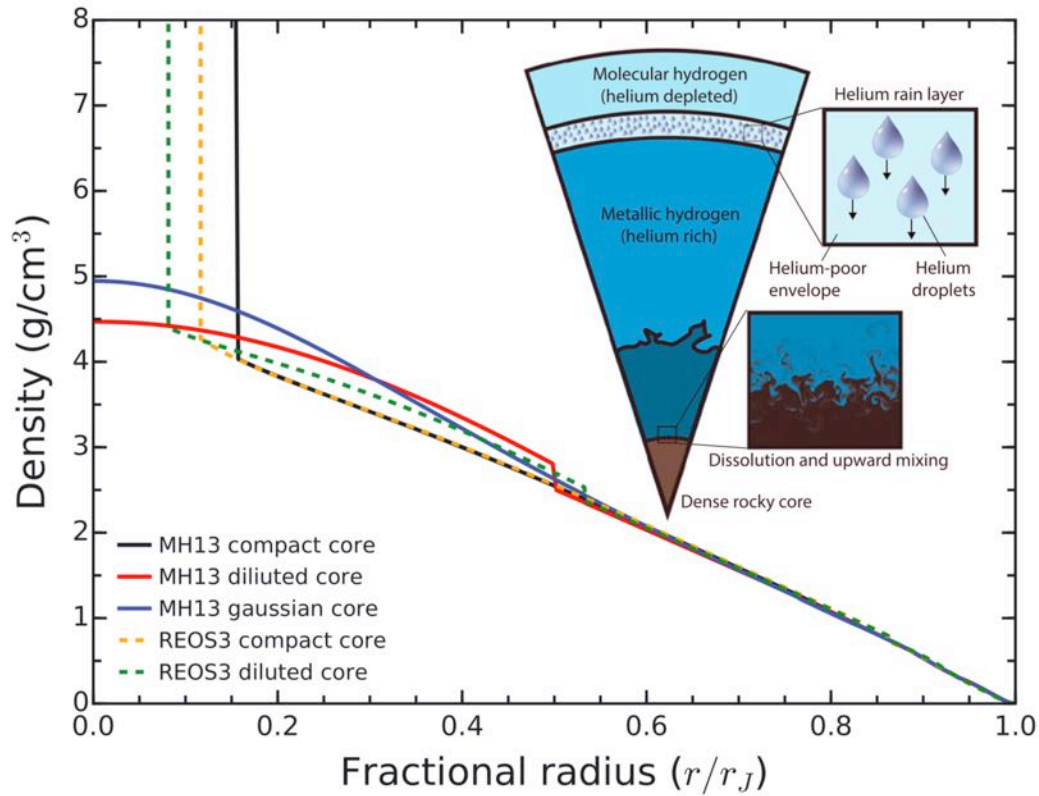
- Elements can separate during the evolution, e.g. helium rain
- Energy transport is usually/often by convection. How do they interact?



Composition effects likely explain some puzzles in the cooling of the Solar System giant planets

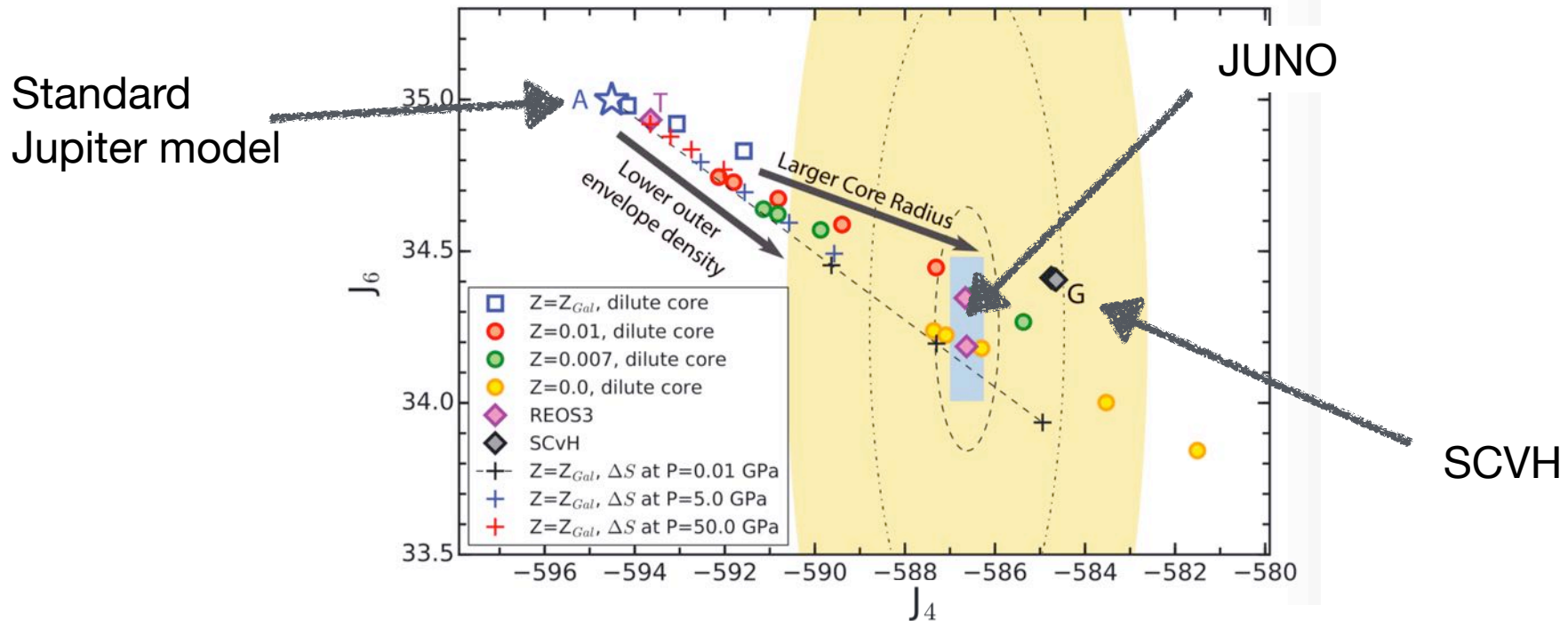


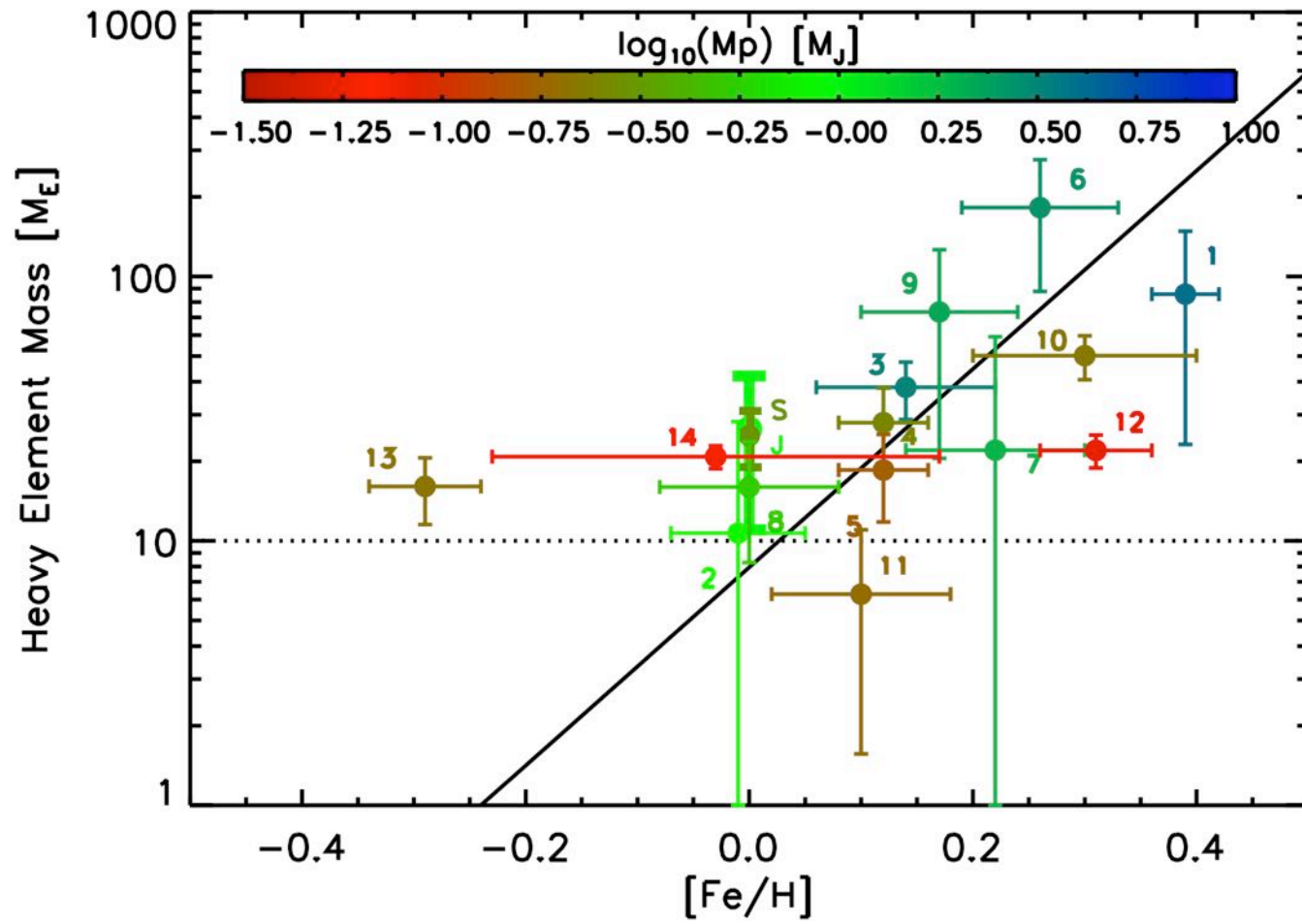
Fortney et al. (2011)



Juno gravitational moments suggest a diluted core in Jupiter

Wahl et al. (2017)





Miller & Fortney (2011)

This talk

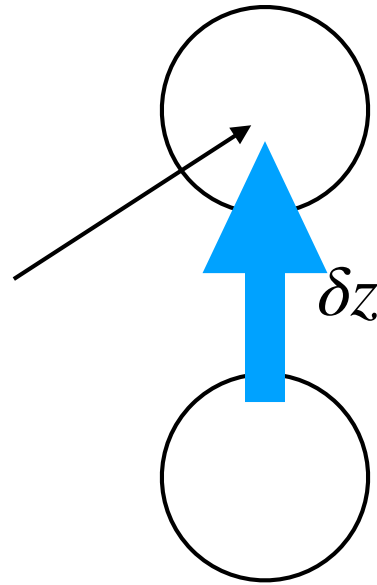
- Interaction of convection with composition gradients and how it might apply to Jupiter
- Convective layers in 1D
- 2D Boussinesq simulations of a salinity gradient cooled from above

Interaction between convection and composition gradients

Thermal convection

Fluid element

$$\frac{\delta\rho}{\rho} = \left(\frac{d \ln P}{dz} \delta z \right)^{1/\gamma}$$



Background

$$\frac{\delta\rho}{\rho} = \frac{d \ln \rho}{dz} \delta z$$

The adiabatic gradient is marginally stable

$$\nabla > \nabla_{\text{ad}} \Rightarrow \text{convection}$$

$$\nabla = \frac{d \ln T}{d \ln P}$$

$$\nabla_{\text{ad}} = \left. \frac{\partial \ln T}{\partial \ln P} \right|_s$$

Interaction between convection and composition gradients

Thermal convection

Large heat flux $F \sim \rho v_c c_p T (\nabla - \nabla_{\text{ad}})$

$$v_c^2 \sim g \ell (\nabla - \nabla_{\text{ad}})$$

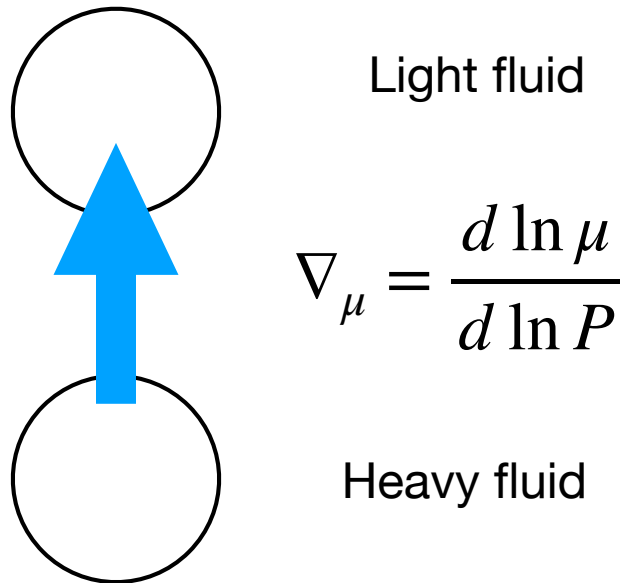
↑
mixing length

Typically need $\nabla - \nabla_{\text{ad}} \lll 1$

Convection \Leftrightarrow **Adiabatic interior**

Interaction between convection and composition gradients

With composition gradients:



Ledoux criterion

$$\nabla > \nabla_{\text{ad}} + \frac{\chi_{\mu}}{\chi_T} \nabla_{\mu}$$

=> convection

There is then a composition flux $F_X \sim \rho v_c \ell \nabla_{\mu}$

Convection

<=>

Uniform composition

What happens when $\nabla_{ad} < \nabla < \nabla_{ad} + \nabla_{\mu}$?

Hot salty

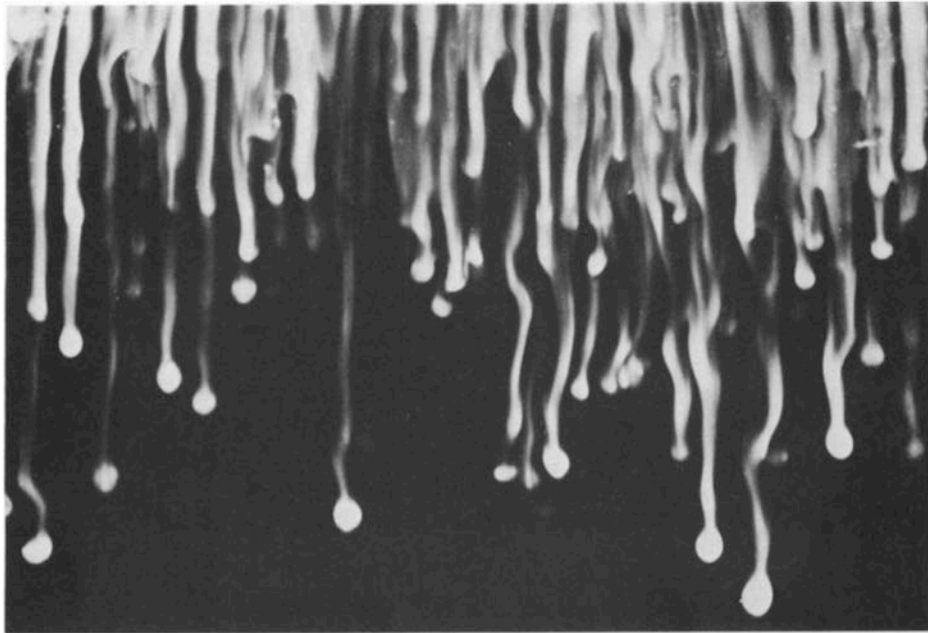


FIGURE 1. A field of salt fingers formed by setting up a stable temperature gradient and pouring a little salt solution on top. The downward-moving fingers were made visible by adding fluorescein to the salt and lighting through a slit from below.

Cold fresh

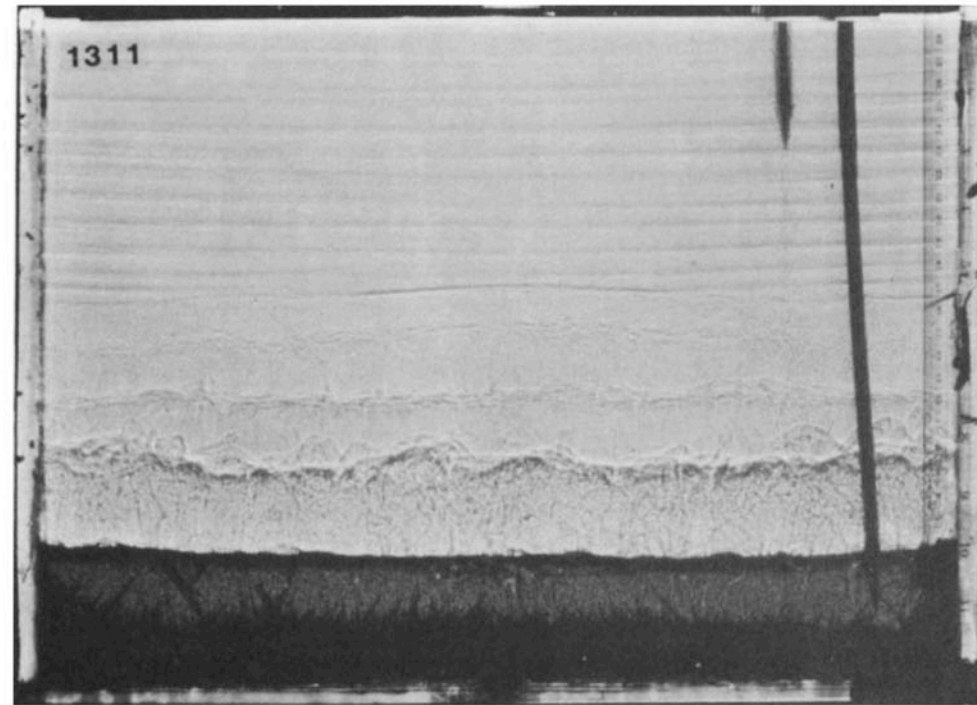


FIGURE 2. A series of convecting layers and 'diffusive' interfaces, formed by heating a gradient of K_2CO_3 solution from below. In this experiment the heating was provided by a hot layer of

Cold fresh

Hot salty

"Fingering regime"

Thermally **stable**

compositionally **unstable**

Thermohaline convection

Thermally **unstable**

compositionally **stable**

Non-linear outcome in both cases: layering and staircases

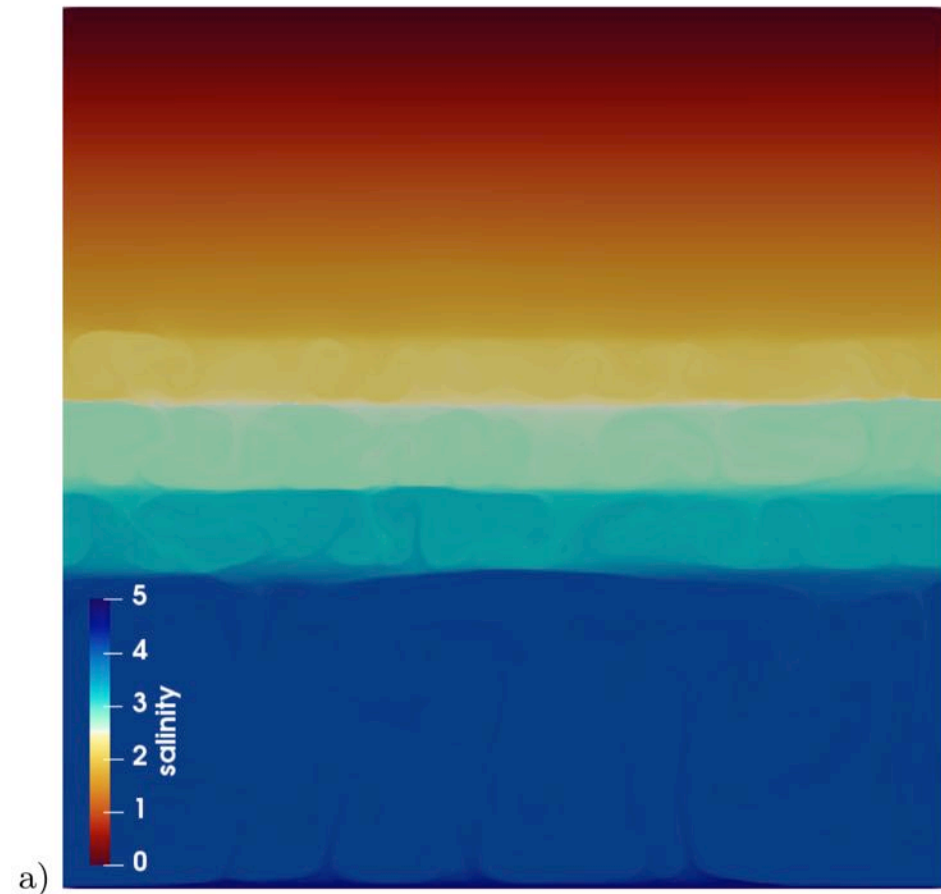
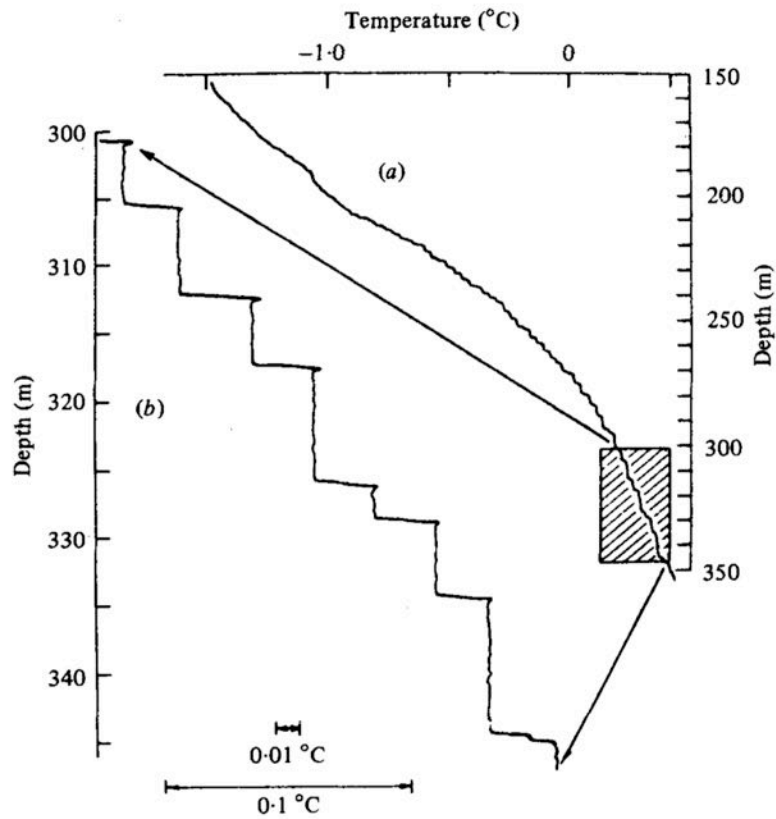


FIGURE 8. A temperature profile obtained under an Arctic ice island by Neal, Neshyba & Denner (1969) showing steps formed by the double-diffusive mechanism. (a) Typical temperature profile section. (b) Section of profile recorded at high gain.

Zaussinger & Kupka (2018)

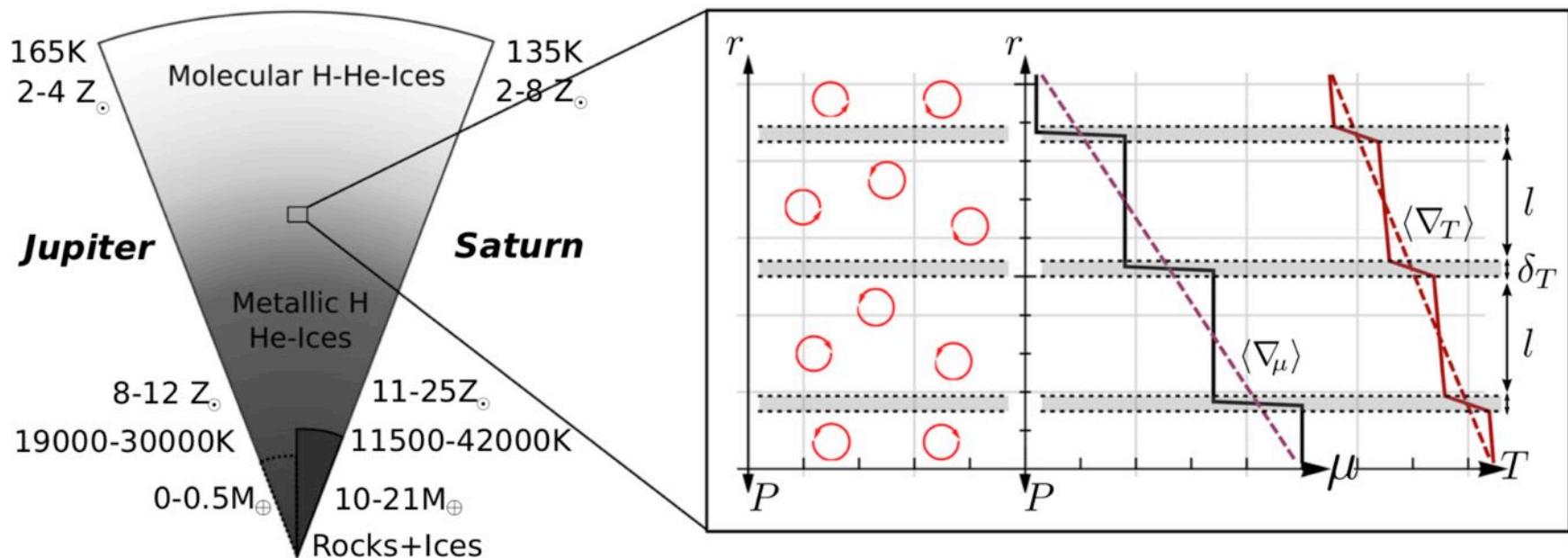
Huppert & Turner (1981)

Layered convection in giant planets

Chabrier & Baraffe (2007): (see also Stevenson 1979,1985)

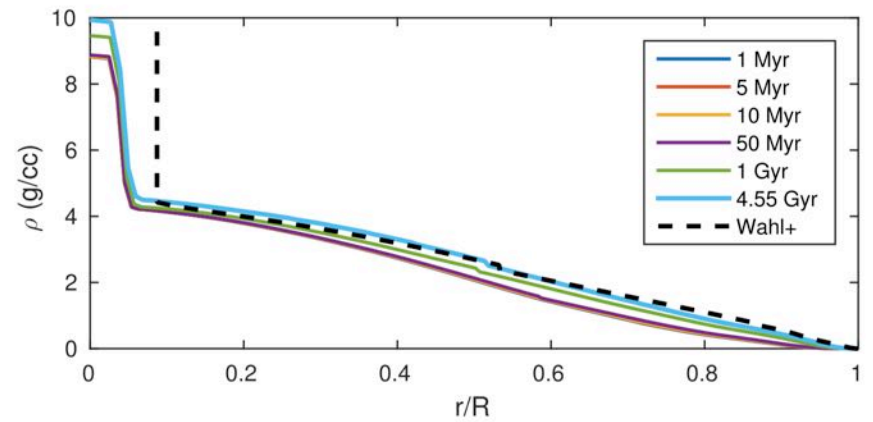
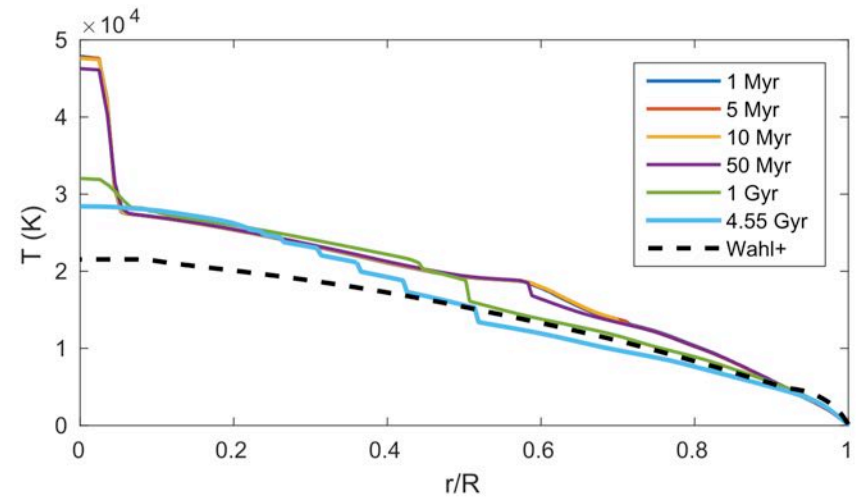
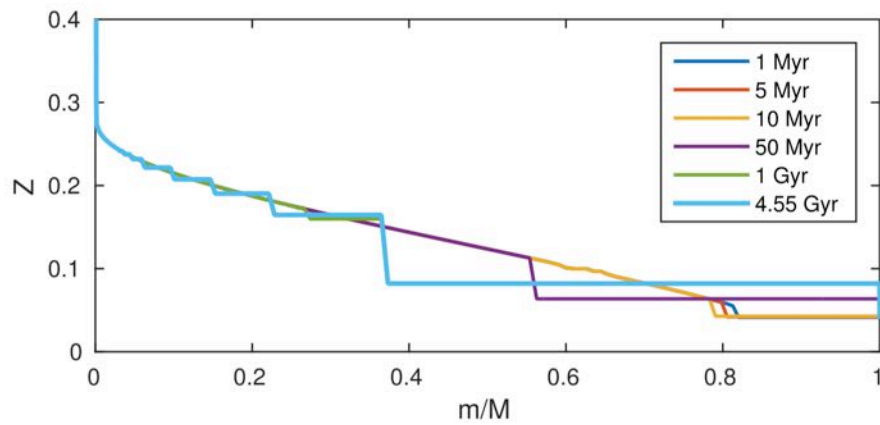
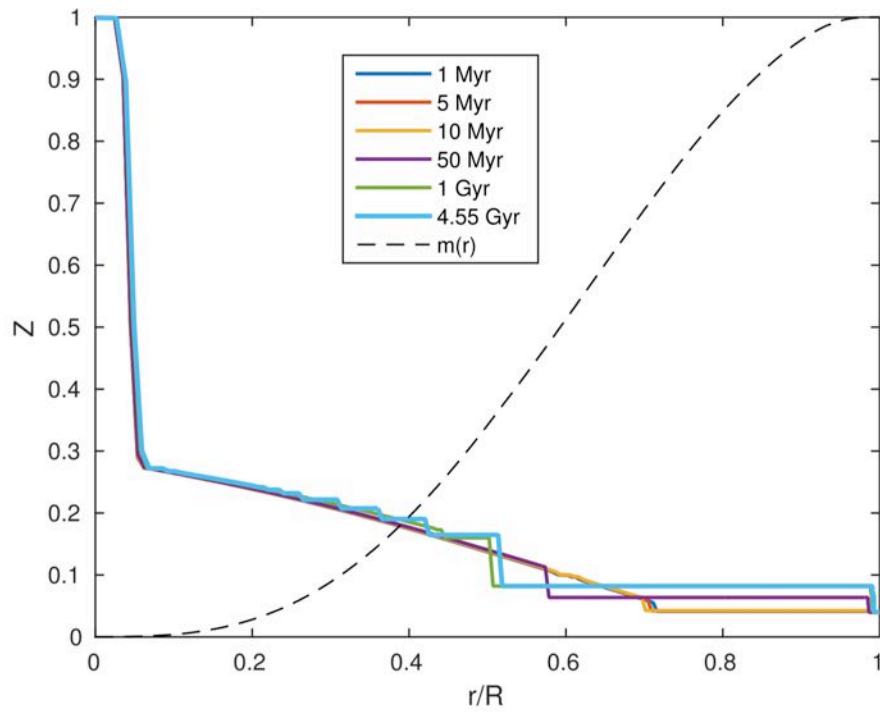
- compositional gradients could inhibit convection in giant planet interiors
- inefficient heat transport could explain inflated hot jupiters?

- estimated layer thickness as $\delta_T \approx (\kappa_T l / \nu)^{1/2} \Rightarrow \sim 10^6$ layers!



Leconte & Chabrier (2012)

Evolution to Jupiter today with heavy elements



Simple model of staircase generation

Solve

$$\frac{\partial T_i}{\partial t} = - \frac{F_{i+1/2} - F_{i-1/2}}{\Delta x}$$
$$\frac{\partial X_i}{\partial t} = - \frac{F_{X,i+1/2} - F_{X,i-1/2}}{\Delta x}$$

on a Cartesian grid with fluxes

$$F_{i+1/2} = a K \nabla_{i+1/2} + a \left(\nabla_{i+1/2} - \nabla_{\text{ad}} \right)$$

$$F_{X,i+1/2} = b \nabla_X$$

convective terms only
turned on when

$$\nabla - \nabla_{\text{ad}} - \nabla_X > 0$$

Gradients (written to be positive
if decreasing outwards)

$$\nabla_{i+1/2} = \frac{T_i - T_{i+1}}{\Delta x}$$

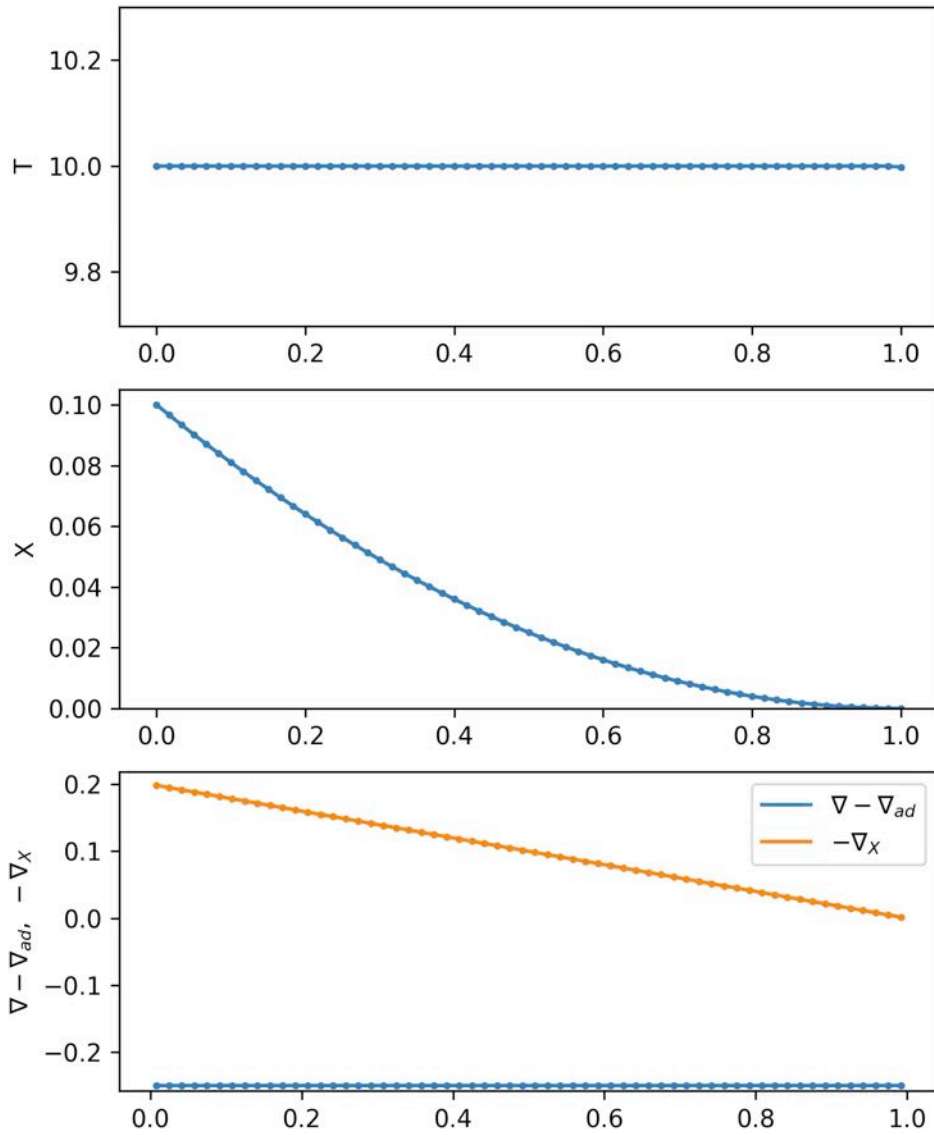
$$\nabla_{X,i+1/2} = \frac{X_i - X_{i+1}}{\Delta x}$$

Show some movies:

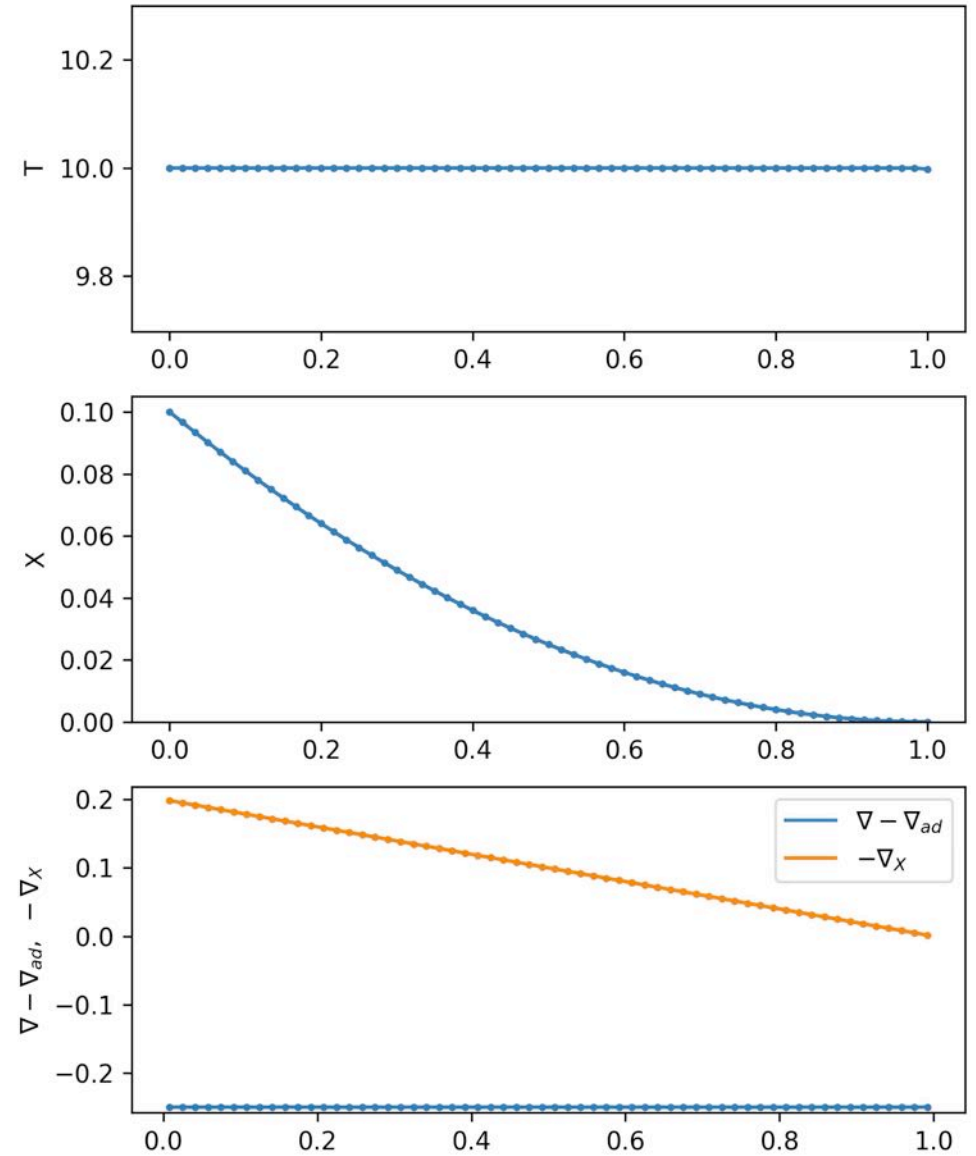
- stairs form only for $K > 0$
- dependence on resolution

Layer formation is a competition between convection and thermal conduction

Schwarzschild, $K=0.00$, $t=0.00$



Ledoux, $K=0.30$, $t=0.00$



(see also discussion in Vazan et al. 2015)

Radko (2005) solved a similar set of equations to study the growth of staircases on a linear background

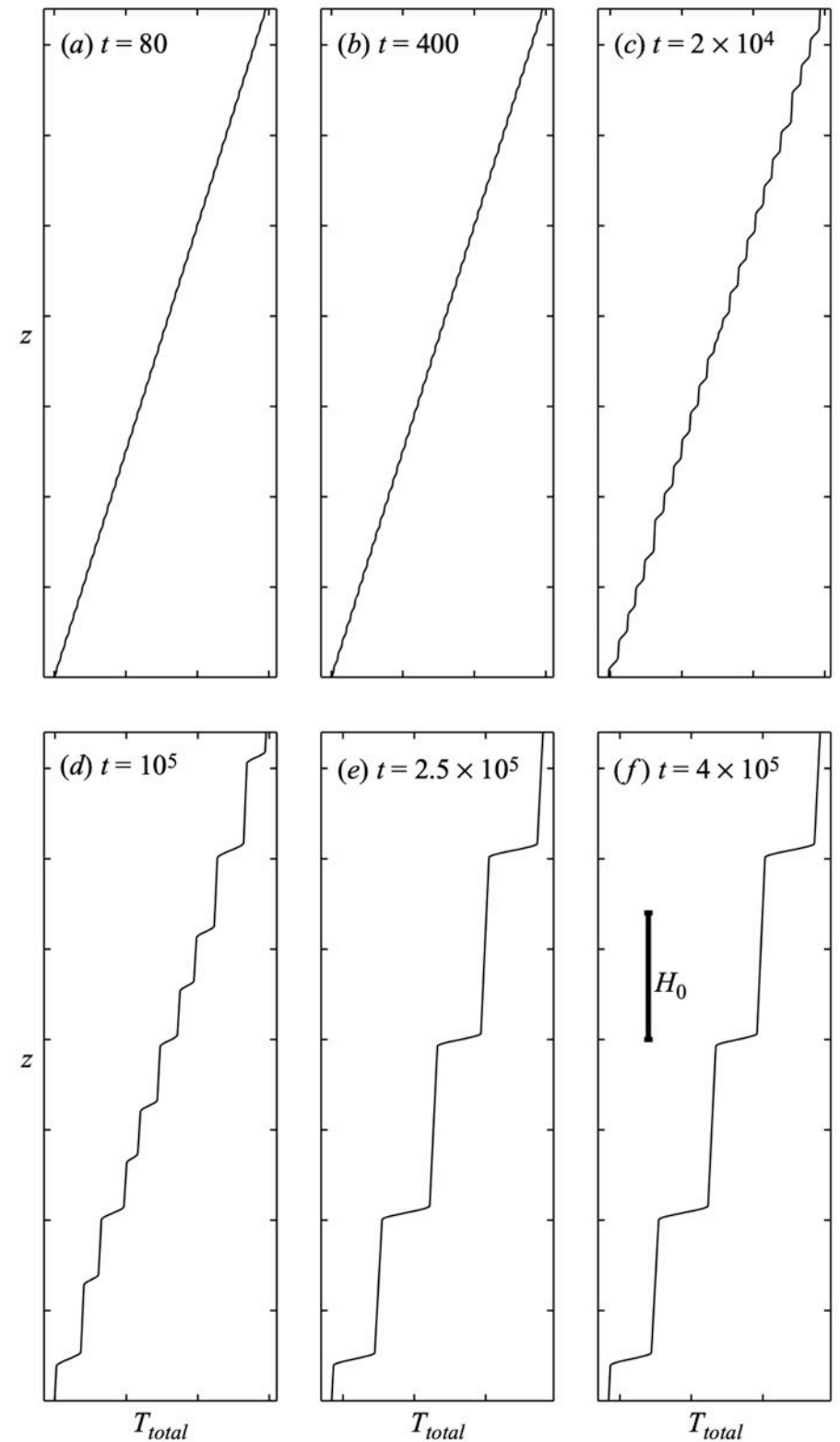
$$\frac{\partial T}{\partial t} = \frac{\partial}{\partial z} F_T - \mu \frac{\partial^4}{\partial z^4} T$$

$$\frac{\partial S}{\partial t} = \frac{\partial}{\partial z} F_S - \mu \frac{\partial^4}{\partial z^4} S$$



Higher order terms stabilize short wavelength unstable modes

Diffusive layer thickness controlled by the parameter μ



Observations of layered convection in the oceans cannot be extrapolated to planets/stars

(see Moll et al. 2016 for discussion)

Important parameters

$$Pr = \frac{\nu}{\kappa_T} \quad \tau = \frac{D}{\kappa_T}$$

Ocean

$$\kappa_T \approx 1.4 \times 10^{-3} \text{ cm}^2/\text{s}$$

$$\nu \approx 10^{-2} \text{ cm}^2/\text{s}$$

$$D \approx 10^{-4} \text{ cm}^2/\text{s}$$

$$Pr \approx 7$$

$$\tau \approx 0.01$$

Planet

$$\kappa_T \approx 10^{-2} - 10^{-1} \text{ cm}^2/\text{s}$$

$$\nu \approx 10^{-3} - 10^{-2} \text{ cm}^2/\text{s}$$

$$D \approx 10^{-4} - 10^{-3} \text{ cm}^2/\text{s}$$

$$Pr \approx 10^{-2} - 1$$

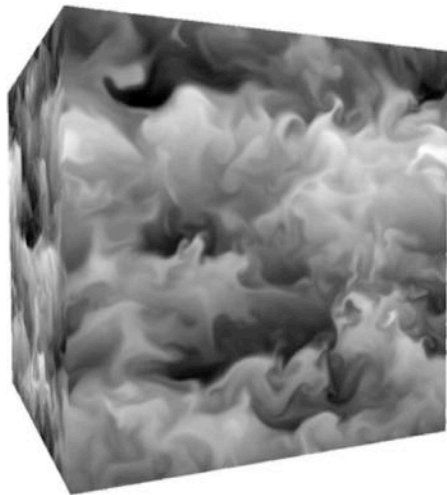
$$\tau \approx 0.01$$

Need numerical simulations at low Pr

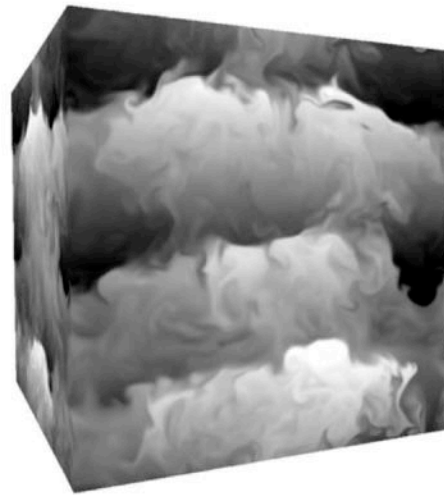
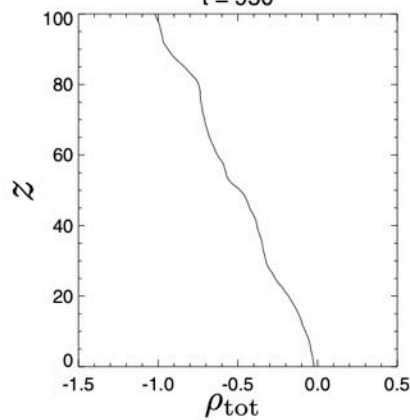
(Stellar conditions are much worse $Pr \ll 1$, $\tau \ll 1$)

Series of numerical studies of semiconvection by Pascale Garaud and collaborators

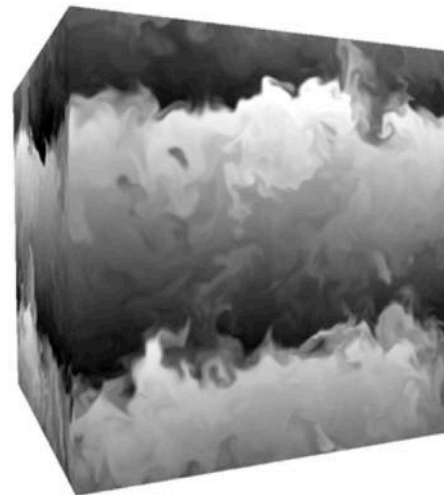
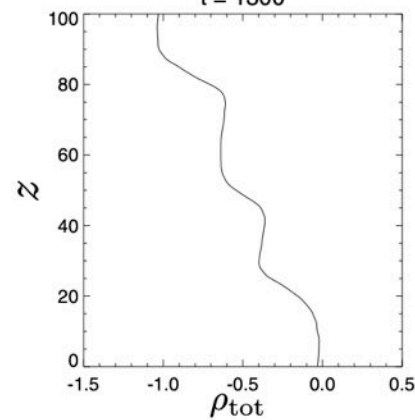
Rosenblum et al. (2011), Mirouh et al. (2012), Wood et al. (2013), Moll et al. (2016)



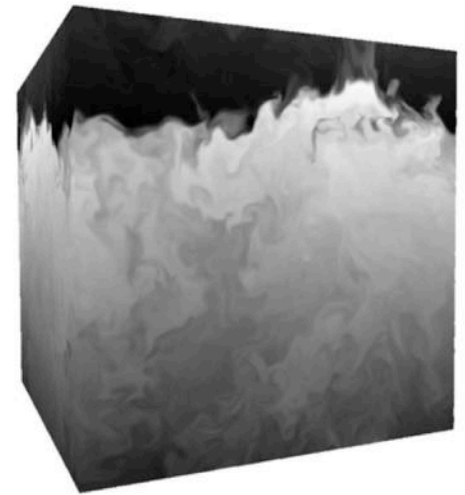
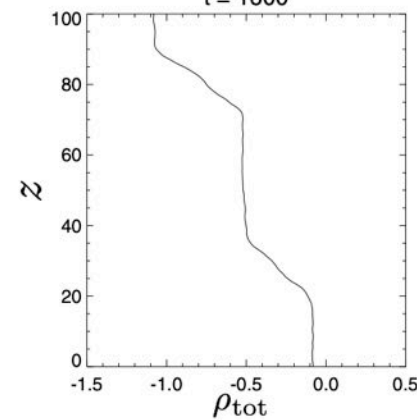
t = 950



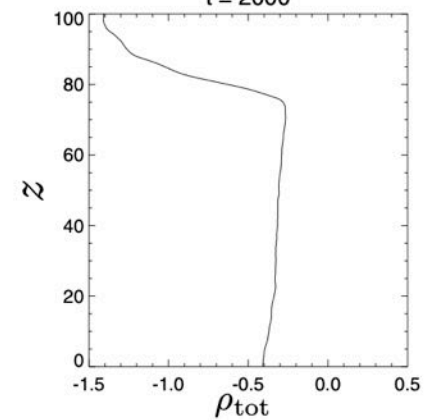
t = 1300



t = 1600

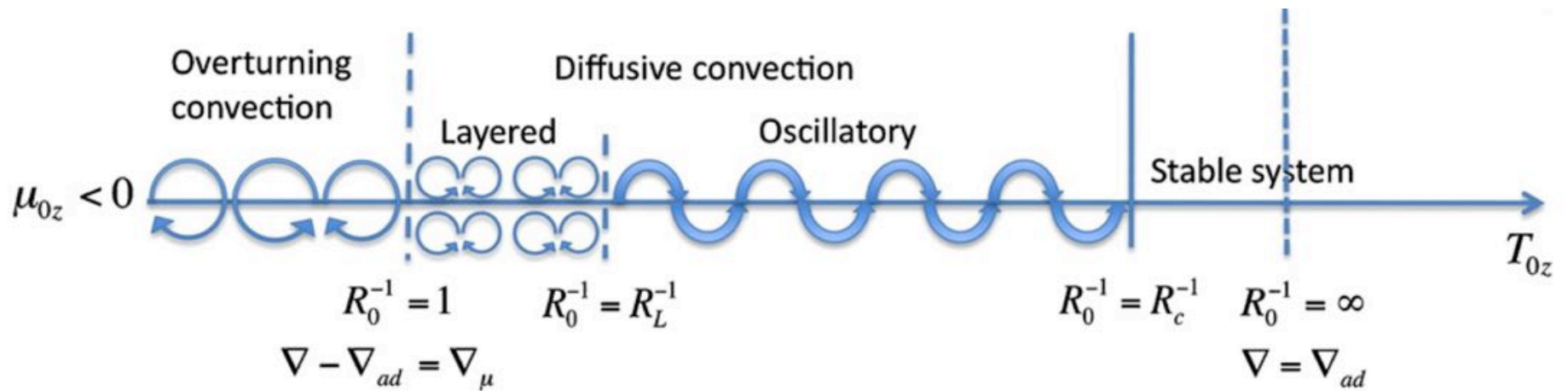


t = 2000



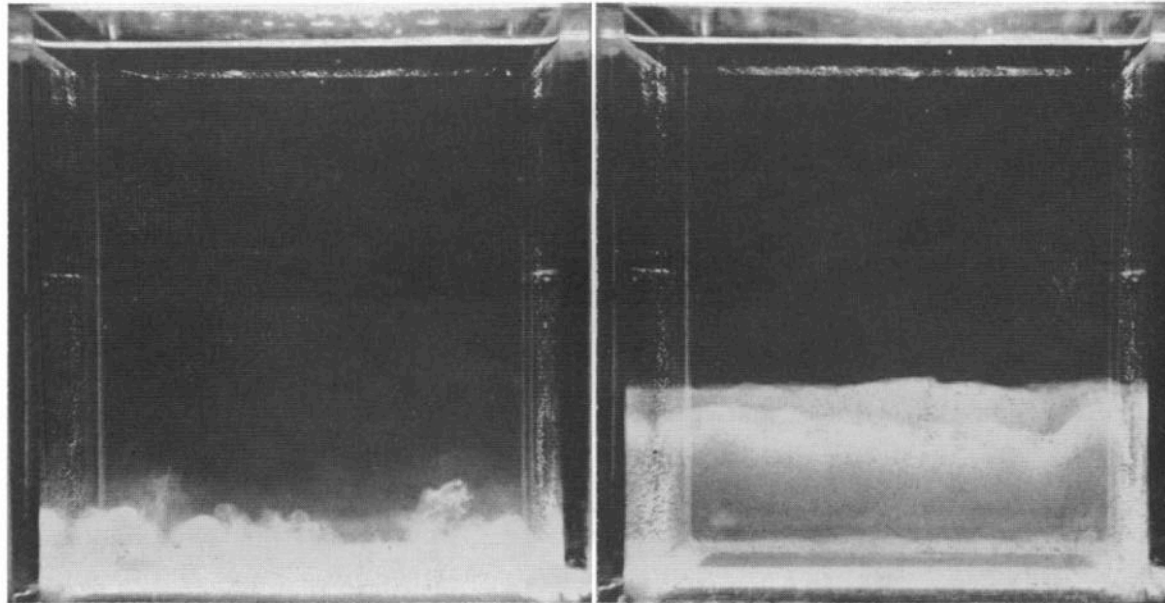
Wood et al. (2013)

Series of numerical studies of semiconvection by Pascale Garaud and collaborators
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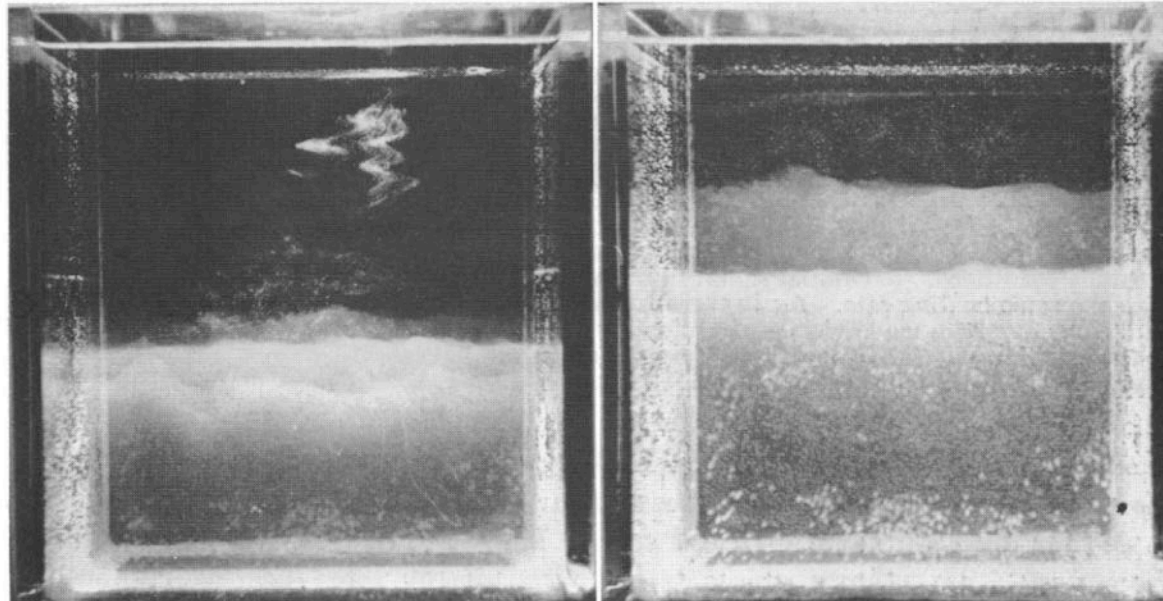
Mirouh et al. (2012)

Turner and Stommel (1964) : A new case of convection in the presence of combined vertical salinity and temperature gradients



a

b



c

d

Convection and mixing in a linear salt gradient cooled from above



Bousinesq approximation

$$\nabla \cdot \mathbf{v}' = 0,$$

$$\frac{\partial T'}{\partial t} = -(\mathbf{v}' \cdot \nabla) T' + \kappa_T \nabla^2 T',$$

$$\frac{\partial X'}{\partial t} = -(\mathbf{v}' \cdot \nabla) X' + \kappa_X \nabla^2 X',$$

$$\frac{\partial \mathbf{v}'}{\partial t} = -(\mathbf{v}' \cdot \nabla) \mathbf{v}' - \frac{\nabla P'}{\rho_b} + \left(\frac{\rho'}{\rho_b} \right) \mathbf{g} + \nu \nabla^2 \mathbf{v}'$$

Rafael Fuentes
(PhD student at McGill)

Equation of state

$$\rho' = \rho_b (\beta X' - \alpha T')$$

$$\alpha = 2.3 \times 10^{-4} \text{ K}^{-1}$$

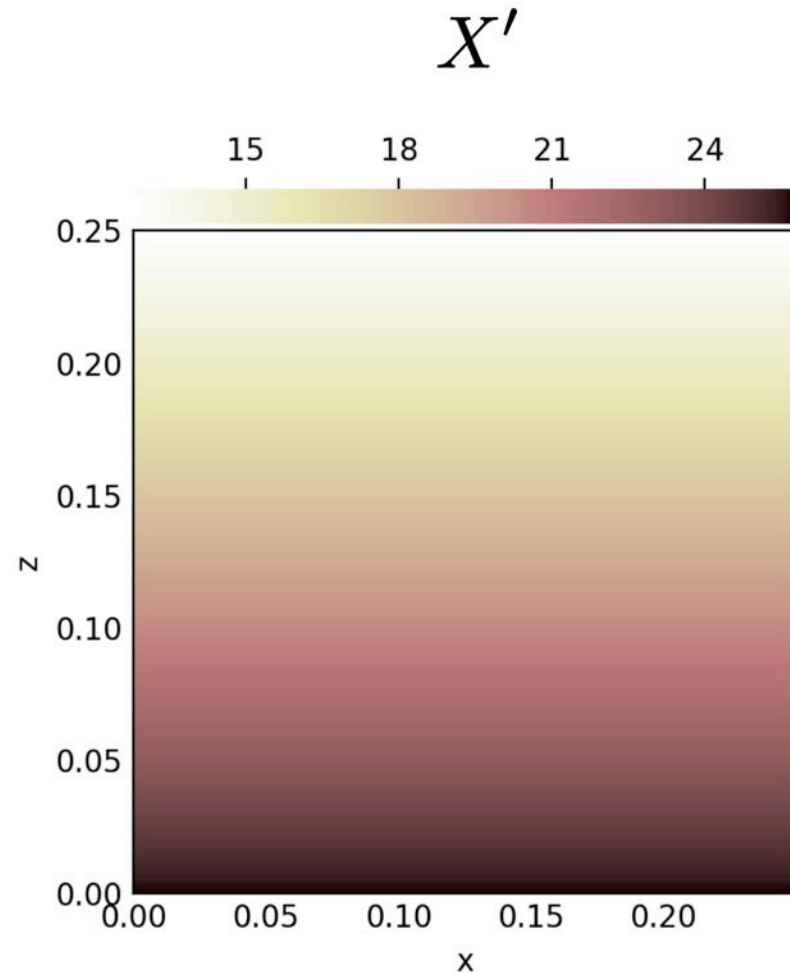
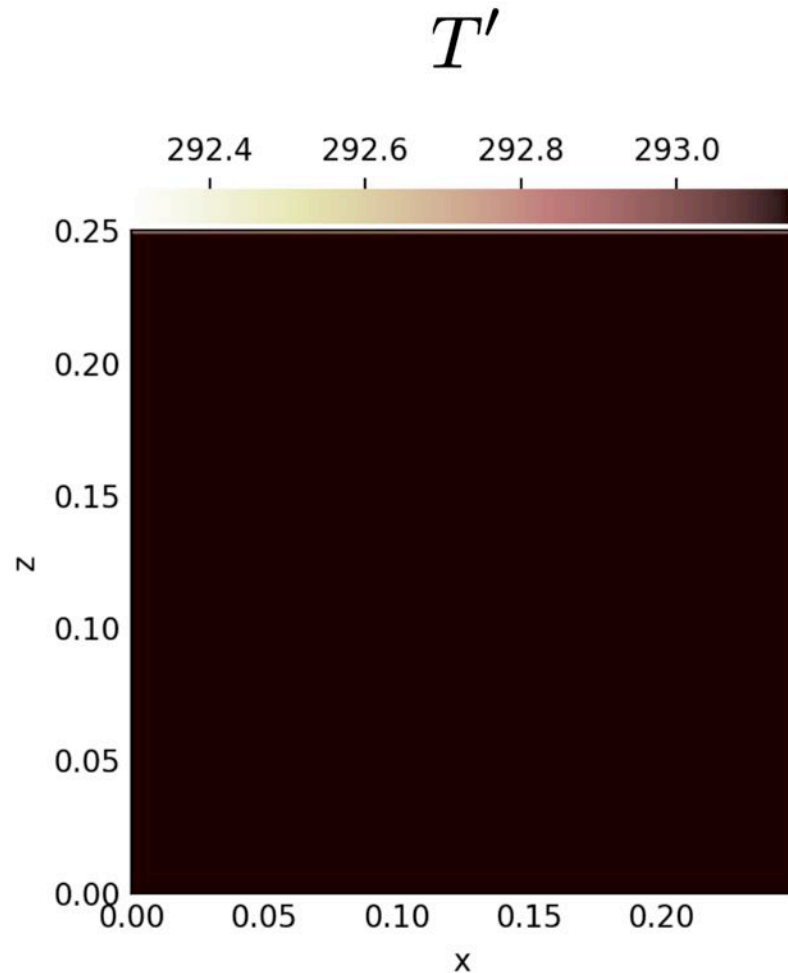
$$\beta = 7.6 \times 10^{-4}$$

Dedalus code <http://dedalus-project.org>

Open source, python based spectral code for solving PDEs

Initial conditions

1. Uniform temperature
2. Linear gradient of solute

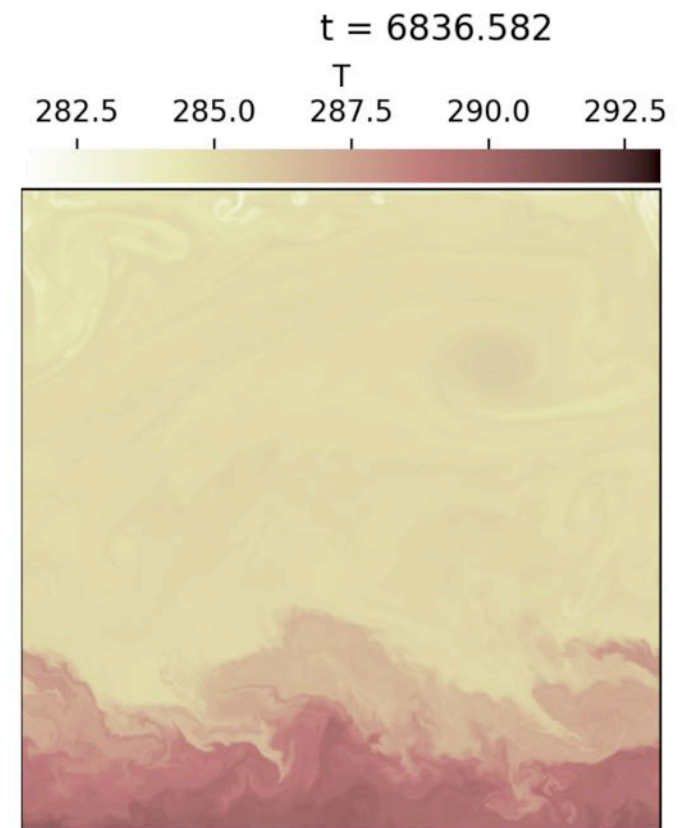
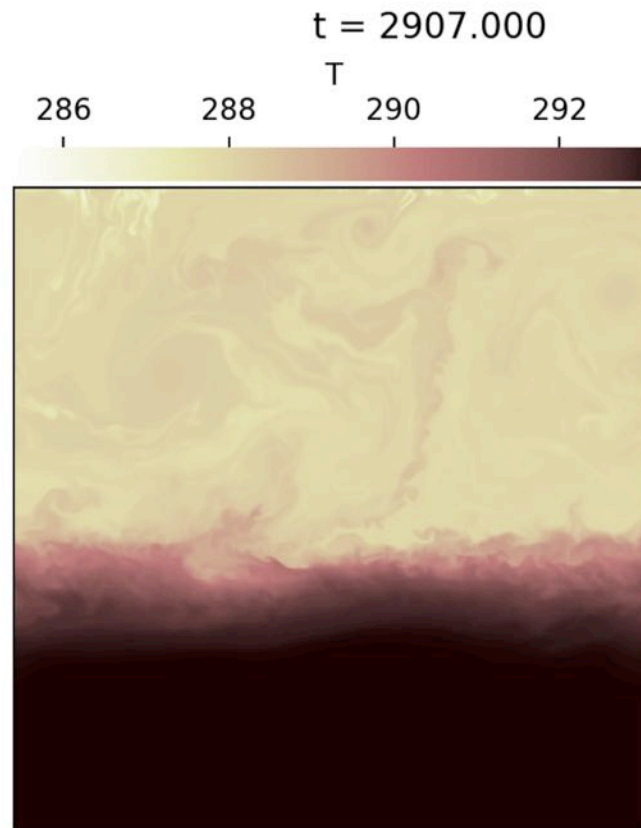
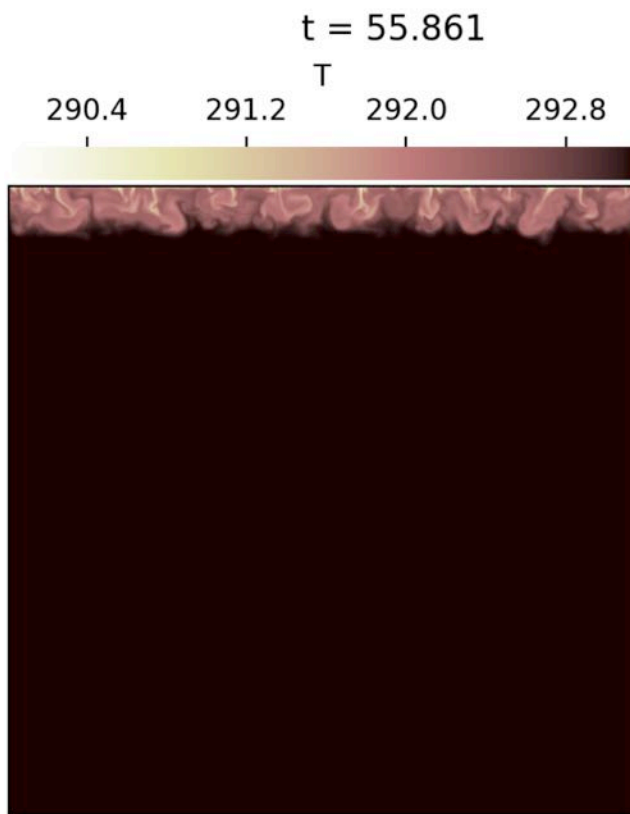


Apply a constant flux at the top of the box, a multiple of the critical flux

$$F_{\text{crit}} = k \frac{\beta}{\alpha} \left| \frac{dX'_0}{dz} \right|$$

Marginally stable temperature gradient

$$Pr = 0.1, \quad \tau = 0.1$$

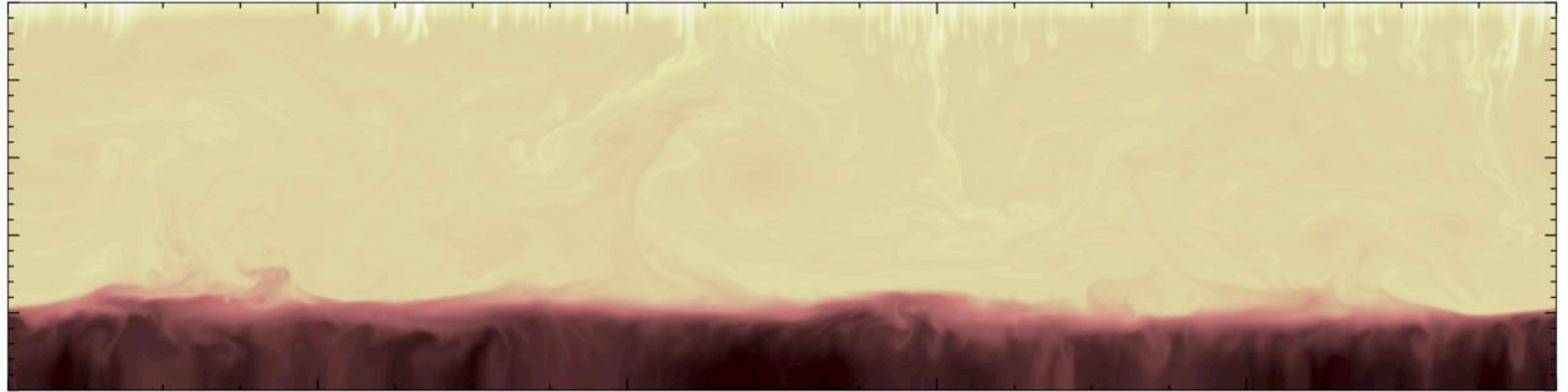
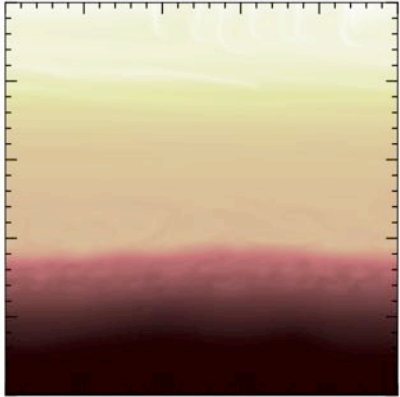


Show some movies:

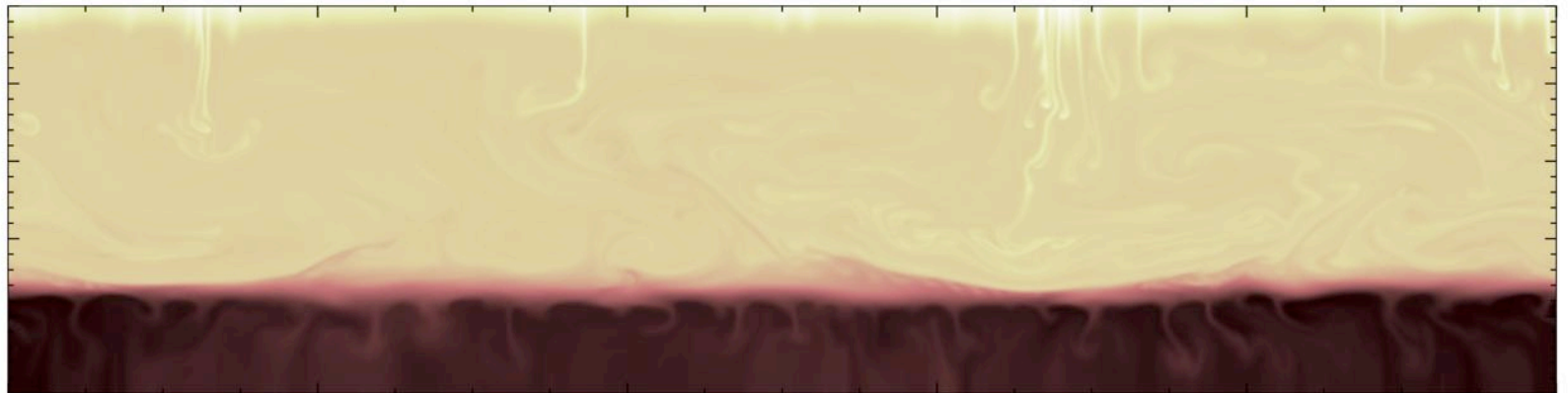
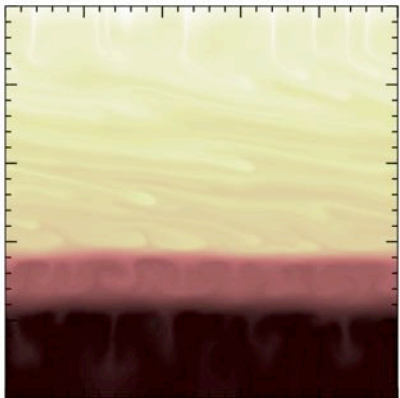
- Shear develops in square boxes
- Goes away with larger aspect ratio
- Differences between $Pr = 7$ and 1

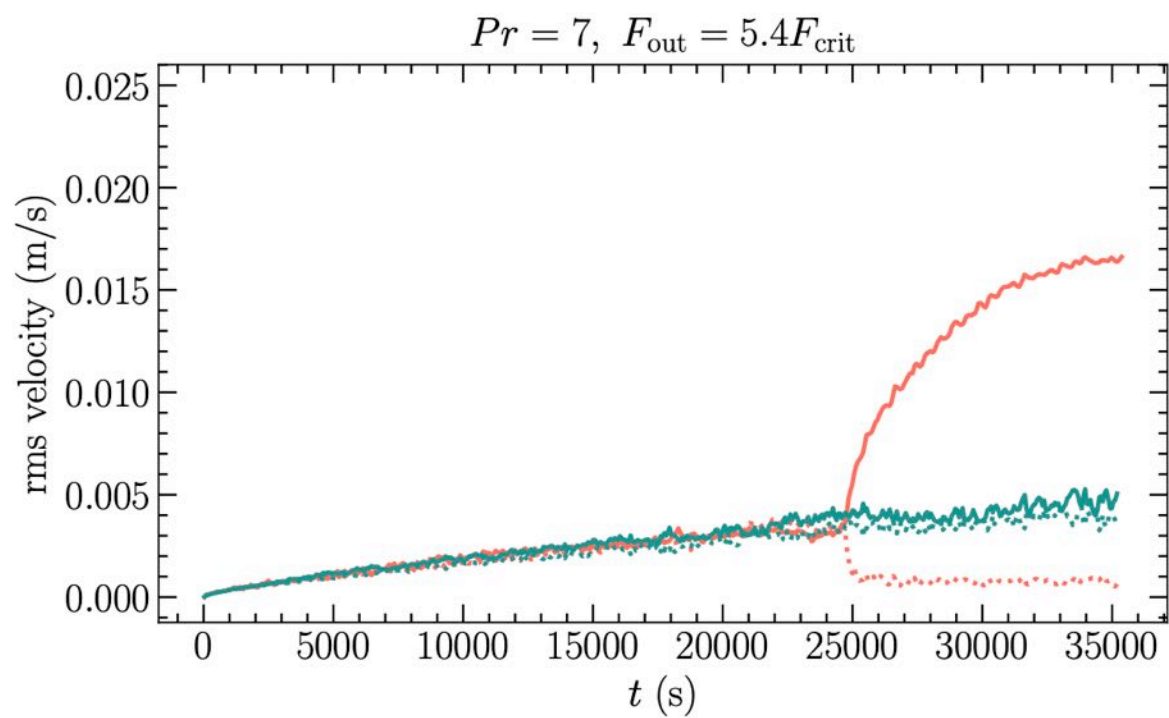
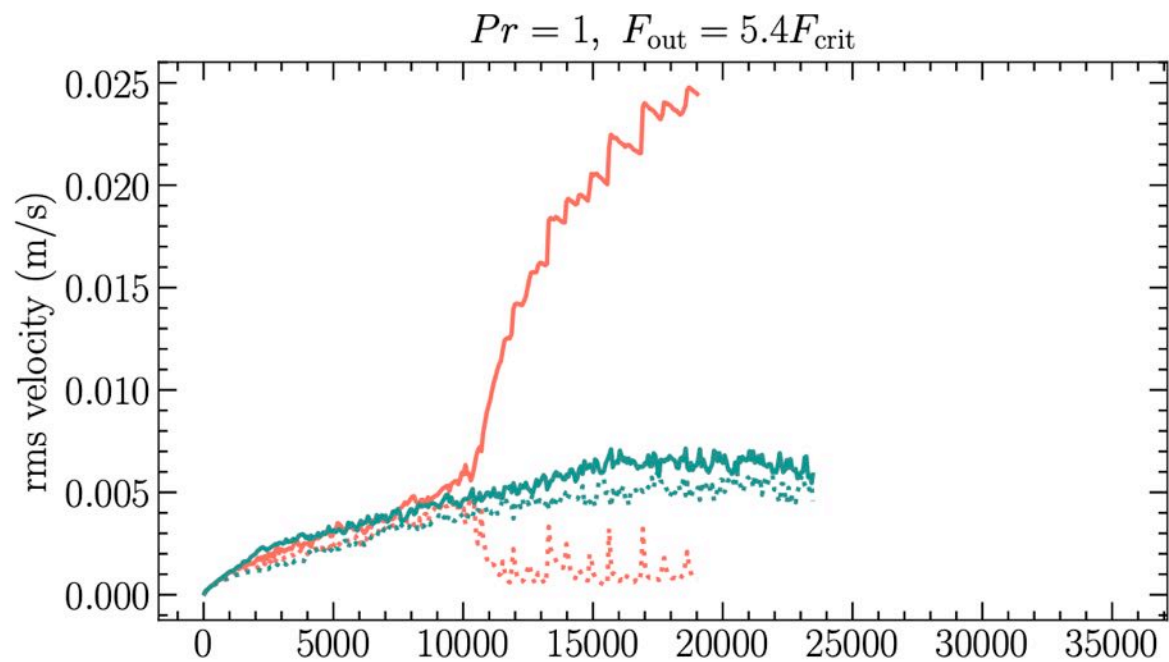
Larger aspect ratio boxes avoid the problem of shear

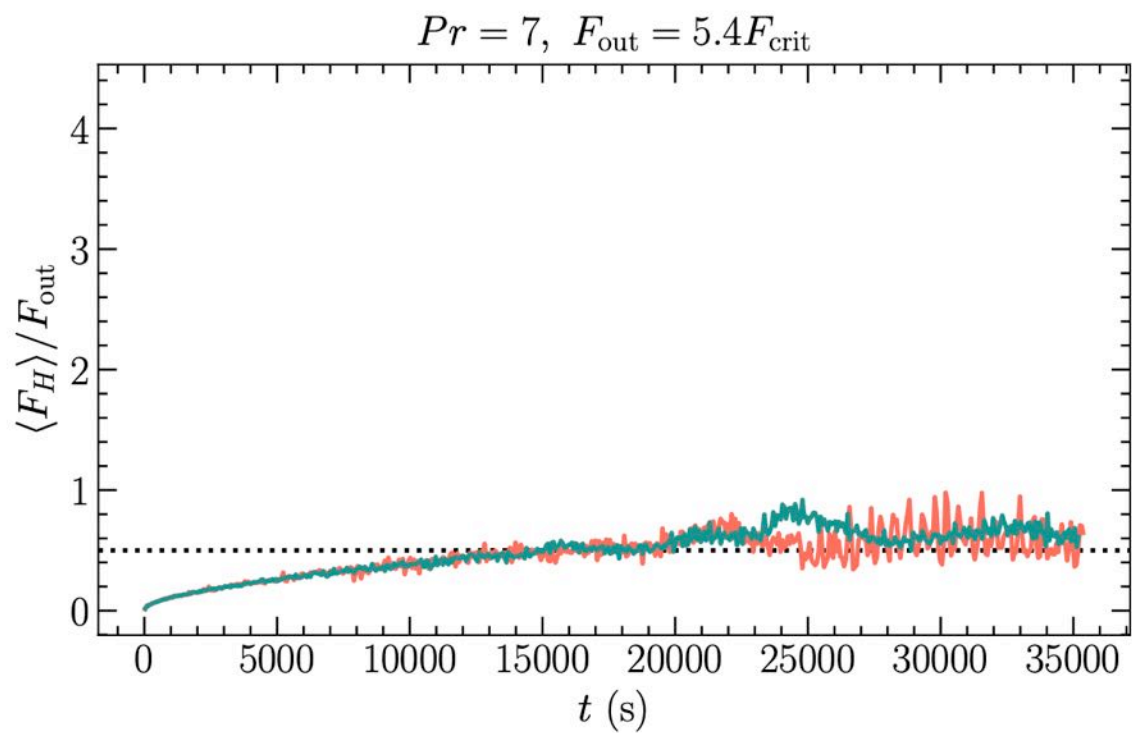
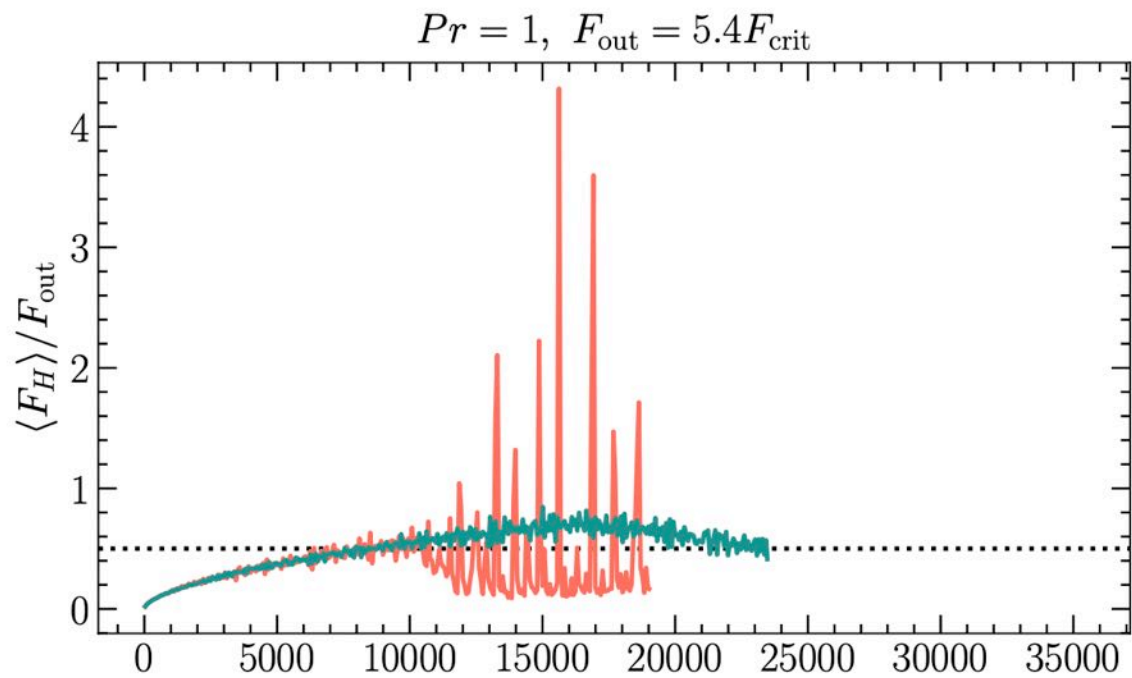
$$Pr = 1, F_{\text{out}} = 5.4F_{\text{crit}}, t = 5.3 \text{ h}$$



$$Pr = 7, F_{\text{out}} = 5.4F_{\text{crit}}, t = 9.8 \text{ h}$$

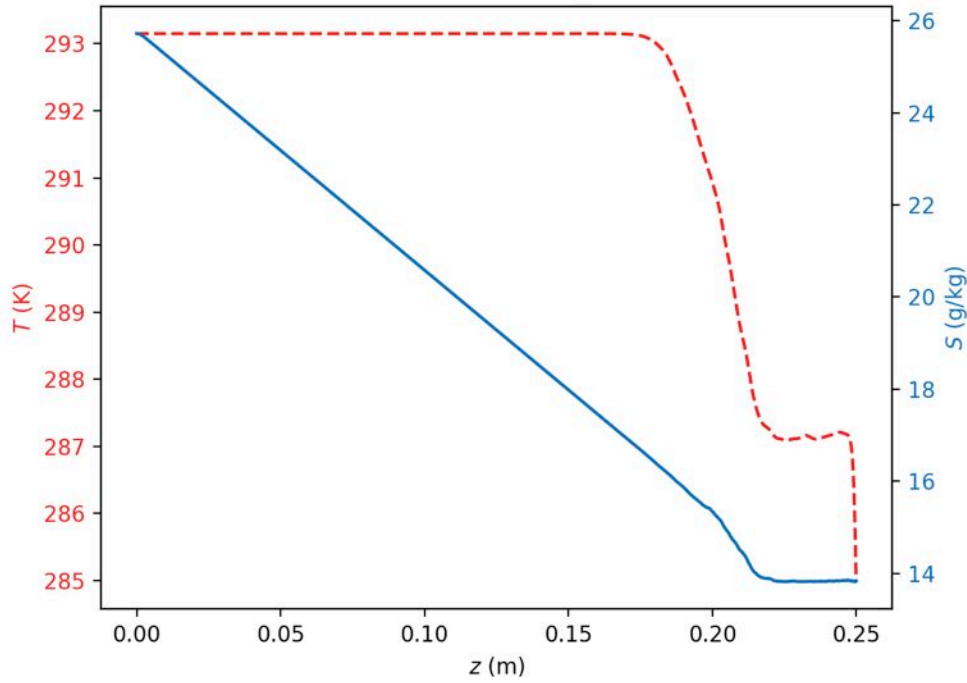




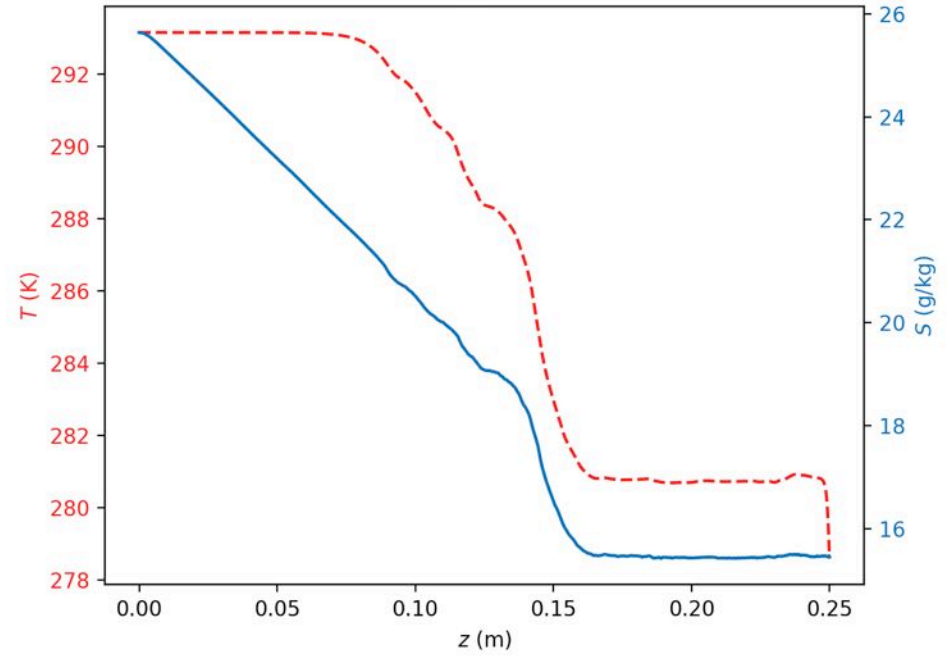


..... $\langle F_T \rangle / F_{\text{out}} = 0.5$ — $L = H$ — $L = 4H$

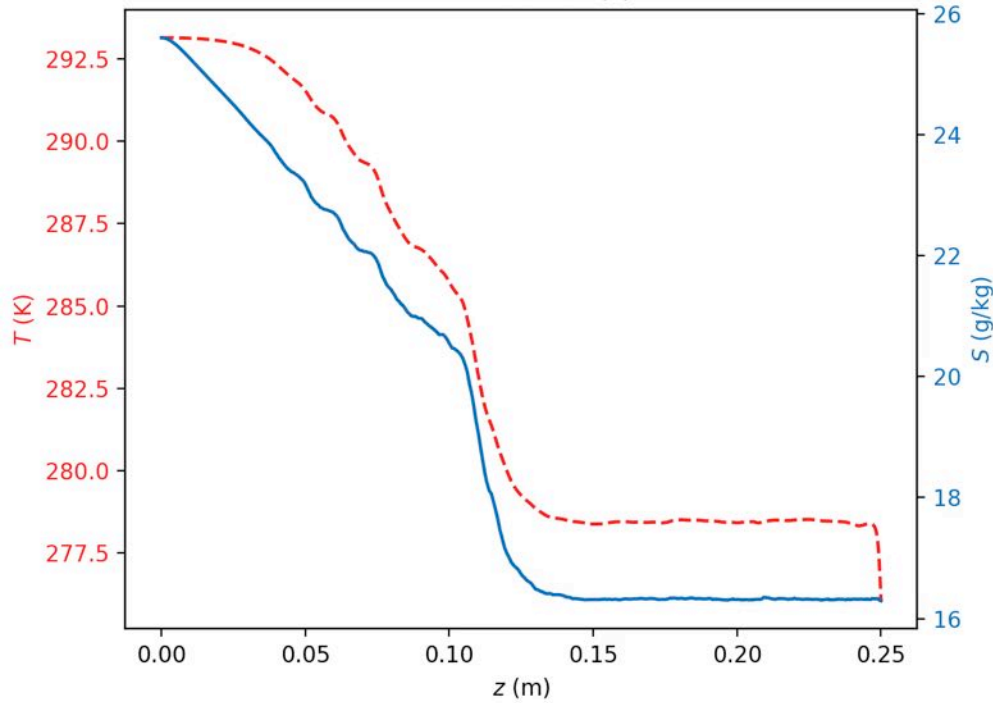
t = 872.100 (s)



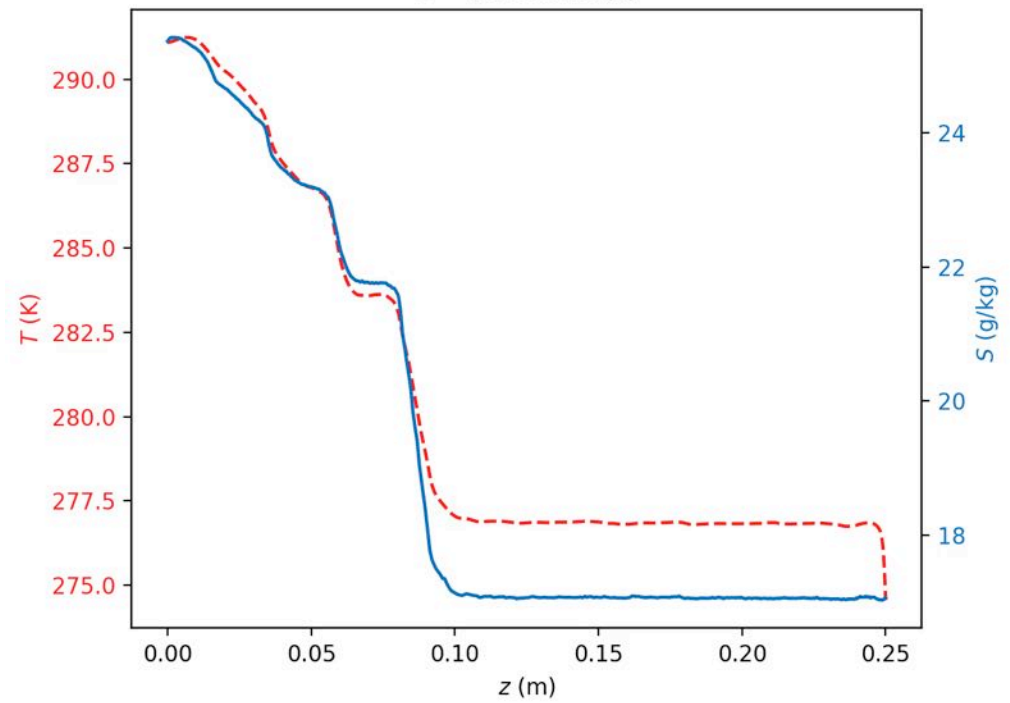
t = 4505.280 (s)



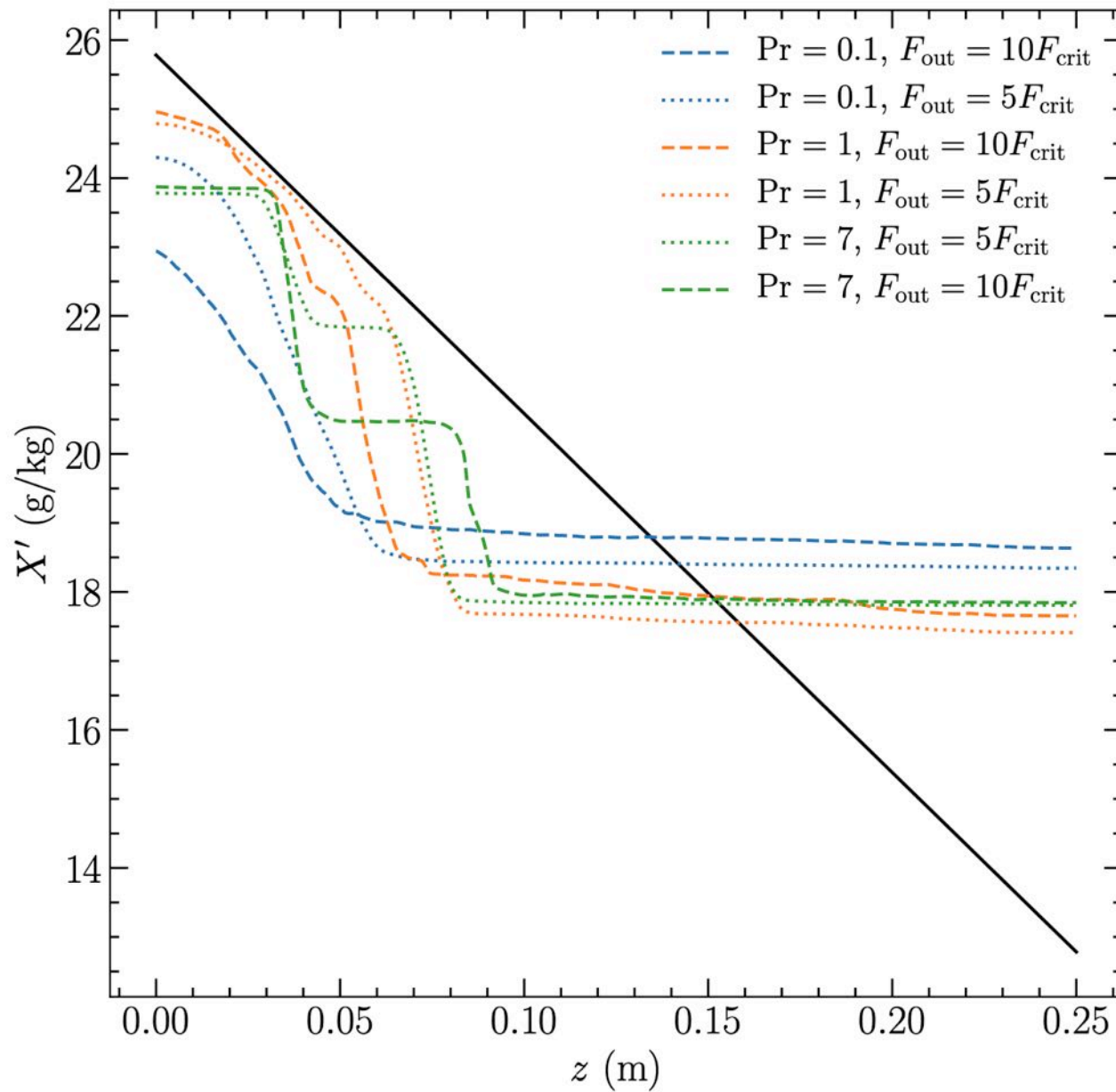
t = 7120.440 (s)



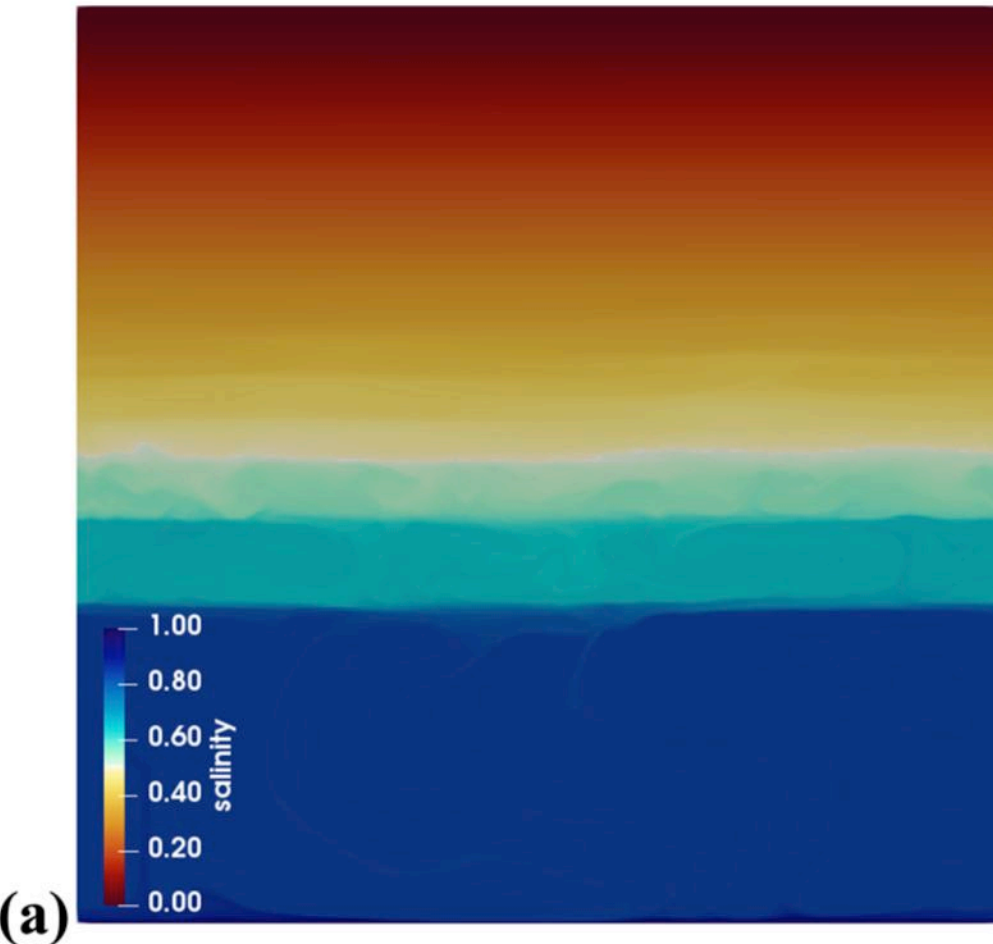
t = 9590.820 (s)



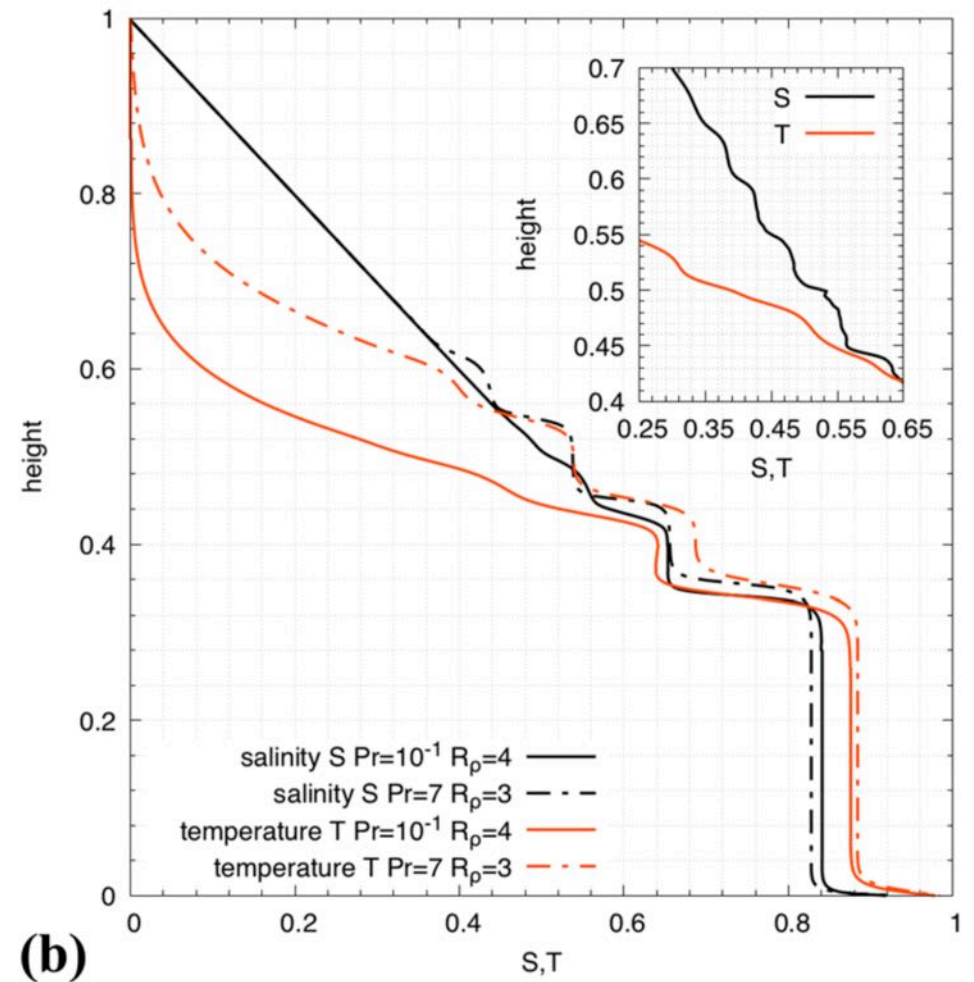
Steps are much less prominent or absent at $Pr < 1$



Layer formation in double-diffusive convection over resting and moving heated plates



$Pr = 0.1$



How quickly does the outer convection zone move inwards?

Turner (1968) argument:

$$\rho_b c_p \Delta T' h = F_{\text{out}} t$$

Energy conservation determines the temperature drop

$$\Delta X' = \frac{1}{2} \left| \frac{dX'_0}{dz} \right| h$$

Solute is mixed throughout the convection zone

When is the temperature difference enough to mix the underlying layer upwards?

$$\alpha \Delta T' = C \beta \Delta X'$$

Two limits: $C = 1$

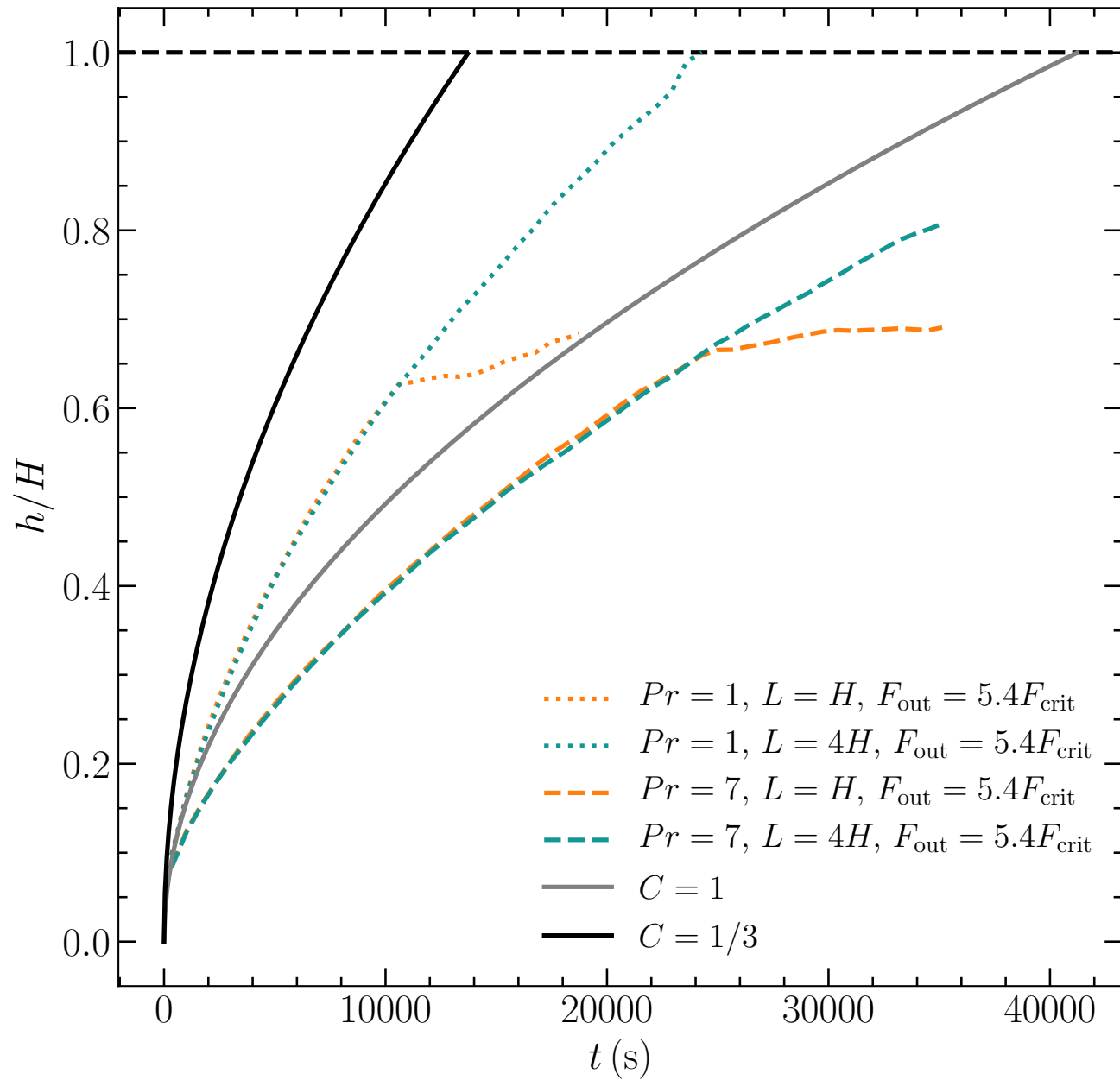
The temperature difference eventually overcomes the composition difference (Ledoux)

$C = 1/3$

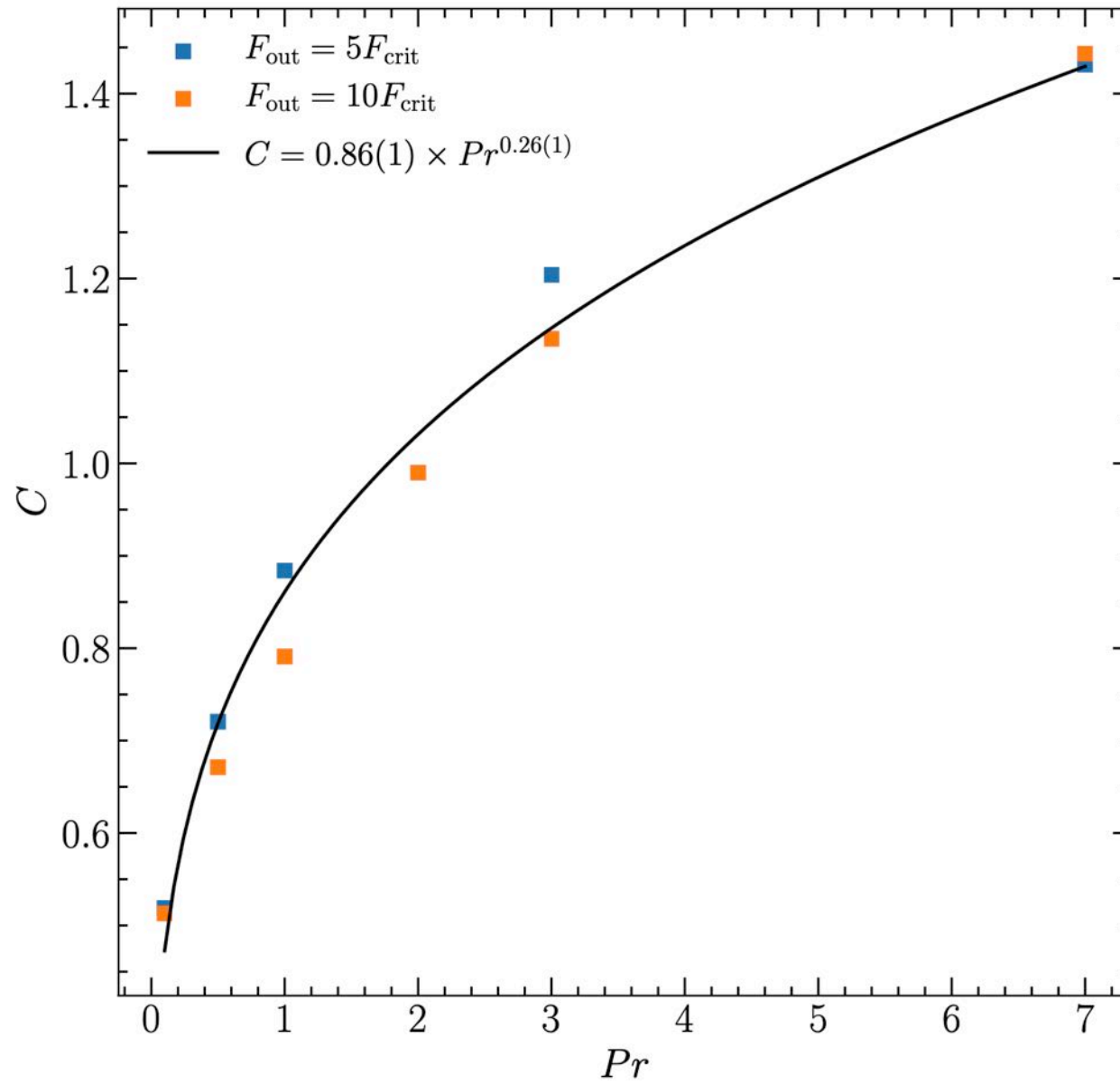
The convection can entrain heavy fluid from below: thermal energy used to lift the material

$$\Rightarrow h(t) = \sqrt{\frac{2}{C}} \left(\frac{F_{\text{out}}}{F_{\text{crit}}} \right)^{1/2} (\kappa_T t)^{1/2}$$

The rate at which the convection zone moves inwards is only weakly-dependent on Pr



The rate at which the convection zone moves inwards is only weakly-dependent on Pr



Conclusions

- Composition gradients should be ubiquitous in planets, and need to be considered in formation, evolution, and structure models
- Convective layers appear in 1D when a composition gradient is heated/cooled
- 1D evolution requires some kind of model for the boundary layer between convection zones in order to resolve them
- This gives another parameter ... analogous to stellar overshoot... how to determine/calibrate it? What is the correct prescription in 1D?
- We are lucky that we can simulate the parameter regime of planets ($Pr \sim 0.01-1$, $\tau \sim 0.01$)
- Aspect ratios >1 avoid generation of shear
- Low $Pr \Rightarrow$ more turbulent flow \Rightarrow more efficient mixing across interfaces
- Role of gravity waves
- Next steps: include stratification (anelastic), rotation